

| | | | | | |
|--|---|--------------|---|---------------|--|
| Pleistocene and Holocene | | Gt | Talus, includes colluvium | QUATERNARY | |
| | | Qal | Alluvial deposits | | |
| Unconformity | | | | | |
| Tradewater Formation | | | | | |
| Ashban | | ms. | top of middle sandstone(s) | PENNSYLVANIAN | |
| | | c. | unnamed coal bed | | |
| | | ls. | top of lower sandstone(s) | | |
| Caseville Formation | | | | | |
| Morrowan | | p. | Pounds Sandstone Member | | |
| | | Br | Battery Rock Sandstone Member | | |
| | | ws. | "Wayside" Member sandstone lentils | | |
| | | Unconformity | | | |
| Ottawaian | | Mk | Kinkaid Limestone | MISSISSIPPIAN | |
| | | Md | Degonia Formation | | |
| | | Mc | Clare Formation | | |
| | | Mpc | Palestine Sandstone and Clare Formation undifferentiated | | |
| | | Mp | Palestine Sandstone | | |
| | | Mm | Menard Limestone | | |
| | | Mw | Waltersburg Formation | | |
| | | Mv | Vienna Limestone | | |
| | | Mt | Tar Springs Sandstone | | |
| | | Mg | Glen Dean Limestone | | |
| | | Mh | Hardinsburg Sandstone | | |
| | | Mha | Haney Limestone (included with "Mcu"/on cross section) | | |
| | | Mcu | Chesterian undifferentiated (includes Valmyeran St. Louis(?) and St. Genevieve(?) Limestones on map) | | |
| | | Mvu | Valmyeran undifferentiated (cross section only) | | |
| | LINE SYMBOLS: Dashed where inferred; dotted where concealed | | | | |
| | Contact | | | | |
| Coal bed | | | | | |
| Fault: bar and ball on downthrown side | | | | | |
| Reverse fault: triangle on upthrown side | | | | | |
| Syncline | | | | | |
| Line of cross section | | | | | |

The Waltersburg Quadrangle is on the southern margin of the Illinois Basin. Three major southwest-trending sets of parallel faults—the Lusk Creek, Raum, and Hobbs Creek Fault Zones—transect the quadrangle. These post-Pennsylvanian faults are part of a northeast-trending fault system in the western portion of the Illinois-Kentucky Fluorspar District. Except in areas near the faults, strata through most of the quadrangle dip 2° to 5° to the northwest.

The Lusk Creek Fault Zone bounds the northwest side of the Dixon Springs Graben. Mississippian strata of the Waltersburg Formation, Menard Limestone, Palestine Sandstone, and Clore Formation on the upthrown northern side are juxtaposed against Pennsylvanian strata of the Tradewater Formation on the downthrown southern side. The fault zone is 1.5 km wide and extends for 1.5 km. The fault is well exposed as much as 1,600 feet wide. The fault zone contains slices of strata ranging from St. Louis? Limestone (Mississippian) to Tradewater Formation (Pennsylvanian). Maximum vertical displacement across the fault zone, estimated from stratigraphic juxtapositions, is about 750 feet. Drilling data, confidential seismic data, and associated drag folds indicate that the fault zone is a normal fault. The fault zone contains horsts within the fault zone expose strata older than the strata on adjacent sides of the fault zone.

The Raum Fault Zone parallels the Lusk Creek and Hobbs Creek Fault Zones and bisects the Dixon Springs Graben. The throw is a few feet to several hundred feet down to the northwest. The fault zone consists of a single fault or two or more subparallel fault slices and generally is narrow, less than 150 feet wide, except where it branches at the southwest corner of the quadrangle. Similarly to the Lusk Creek Fault Zone, small horsts occur within the Raum Fault Zone. Subsurface data, seismic data, and associated drag folds indicate the fault plane dips to the northwest.

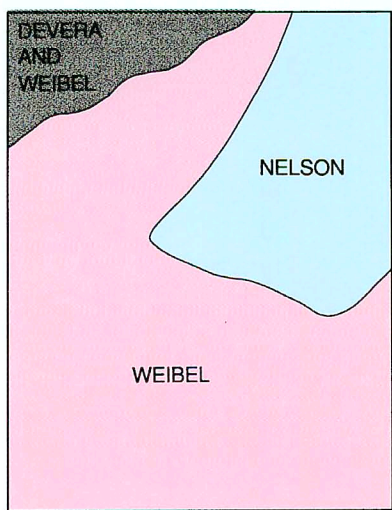
The Dixon Springs Graben is bounded on the southeast side by the Hobbs Creek Fault Zone. The fault zone generally consists of two subparallel faults that range from 3,600 to 6,650 feet apart. The fault block between these bounding faults is offset by subparallel faults forming small grabens within the zone. Rocks within the fault block are similar in age to the strata on the upthrown northwest side. Greater displacement occurs along the southeast bounding fault. Maximum vertical displacement estimated from stratigraphic displacements, is about 250 feet across the fault zone. Seismic data support the field mapping interpretation that most faults in this zone are normal.

We interpret the surface faults to be the result of two episodes of deformation. The first episode, which probably occurred during the latest Pennsylvanian or Permian, was compressional and moved hanging wall strata upward along the Lusk Creek and Raum Fault Zones. During the second episode, which was extensional and probably occurred during the early Mesozoic, strata on the hanging wall moved downward, generally along a different fault plane. The narrow horsts in the Lusk Creek and Raum Fault Zones are slices of rock uplifted during the compressional episode but not downthrown during the extensional episode.

A broad, discontinuous, asymmetric syncline occurs between the Lusk Creek and Raum Fault Zones. The axis of the syncline is adjacent to the Lusk Creek Fault Zone. Dips on the northwest flank are steeper than dips on the southeast flank. We interpret the syncline as a drag fold that formed during the extensional episode of deformation.

Joints in Pennsylvanian sandstones in the quadrangle are generally oriented either parallel or perpendicular to fault traces. Locally joint systems appear to control the orientation of valleys and ravines within the Dixon Springs Graben.

The principal geologic resources in the Waltersburg Quadrangle are limestone for construction and agriculture and sandstone for construction. Small deposits of fluorite and related minerals have been mined locally, suggesting that additional resources may be present in the quadrangle. Resources



C. Pius Weibel, W. John Nelson and Joseph A. Devera
1991