STATE OF ILLINOIS DEPARTMENT OF REGISTRATION AND EDUCATION



STRATIGRAPHY AND MINERALOGY OF THE WISCONSINAN LOESSES OF ILLINOIS

John C. Frye H. D. Glass H. B. Willman

ILLINOIS STATE GEOLOGICAL SURVEY
John C. Frye, Chief
URBANA
CIRCULAR 334
1962



STRATIGRAPHY AND MINERALOGY OF THE WISCONSINAN LOESSES OF ILLINOIS

John C. Frye, H. D. Glass, and H. B. Willman

ABSTRACT

Loess deposits of Wisconsinan age cover more than three-fourths of Illinois and are important as ceramic raw materials, as parent materials of soil, in construction, and in the study of ground-water geology. This report contains the results of analyses by X-ray diffraction, petrographic microscopy, or wet chemistry of 583 samples of loess and related materials. The clay mineral and heavy mineral assemblages of the several stratigraphic units are significantly different and are characteristic of the units. The differences in mineral content also indicate from which major outwash-carrying valleys the several loesses originated and relate the loesses to the up-valley glaciers that provided the outwash. The striking decrease from west to east (Mississippi-Illinois Valley to Ohio-Wabash Valley) in the relative abundance of montmorillonite is consistent with the changes in the mineralogy of the tills.

INTRODUCTION

Loess of Wisconsinan age covers more than three-fourths of Illinois (Smith, 1942, fig. 3; Leighton and Willman, 1950, fig. 4). It ranges from a thin film to more than 50 feet thick and in much of the region it directly overlies glacial drift. Loess and the associated glacial drift compose the prime parent materials of Illinois soils, they constitute a widespread ceramic raw material, they exert a significant control on our ground-water supplies, and they are an important factor in the majority of the construction projects in the state. An adequate knowledge of the stratigraphy and mineral composition of these deposits obviously is needed. The State Geological Survey has been making a mineralogical analysis of these deposits, primarily of the clay-, silt-, and sand-size fractions that volumetrically compose most of their bulk. The results of these studies are reported in two Circulars—this Circular describing the studies of the loesses and

outwash, and a subsequent one that will report data from the glacial tills.

For many years the origin and stratigraphy of the loess deposits of the Upper Mississippi Valley region have been subjects of interest (Udden, 1894, 1898; Chamberlin, 1897; Leverett, 1899; Savage, 1915; Leighton, 1917, 1931; Smith, 1942; Kay and Graham, 1943; Leighton and Willman, 1950; Horberg, 1953, 1956; Frye and Willman, 1960; Leonard and Frye, 1960), but little attention has been given to the mineralogy of the deposits. The present study is based on a representative sampling of loesses, tills, and outwash deposits throughout Illinois. The samples were studied by X-ray analyses of the less than 2-micron fraction, by microscopic determination of the light and the heavy mineral fractions within the very fine and fine sand, using methods previously used (Frye, Willman, and Glass, 1960), and by chemical analyses of the gross sample. The locations and stratigraphic positions of the loess samples used, results of X-ray analyses, light and heavy mineral determinations, and chemical analyses are given in tables 1 through 6 at the end of this report.

The X-ray analyses were made by Glass, the heavy mineral preparation and counts were made largely by Constantine Manos, the light mineral counts were made by James Bloom, and the chemical analyses were made by L. D. McVicker. We express our thanks to George E. Ekblaw and Paul R. Shaffer for helpful suggestions during the progress of the work. The radiocarbon dates used in this report were determined in the Washington laboratories of the United States Geological Survey.

STRATIGRAPHY

The classification of the Wisconsinan Stage and the regional stratigraphy and molluscan paleontology of the loesses in Illinois have been described (Frye and Willman, 1960; Leonard and Frye, 1960). Details of the loess stratigraphy are given in the 15 measured geologic sections published in the two earlier Circulars, in the 15 additional sections included with this report, and by the 29 graphic sections shown on plate 1.

The time-stratigraphic units recognized in the Wisconsinan Stage of Illinois include the Altonian Substage (60,000 \pm to 28,000 radiocarbon years B.P.), the Farmdalian Substage (28,000 to 22,000 B.P.), the Woodfordian Substage (22,000 to 12,500 B.P.), the Twocreekan Substage (12,500 to 11,000 B.P.), and the Valderan Substage (11,000 to 5,000 B.P.).

Deposits of Altonian age recently have been classified in Ontario by Dreimanis (1960, 1961; also, deVries and Dreimanis, 1960) into the Dunwich Drift (at the base), the Port Talbot Interstadial bed, and the Southwold Drift. The Southwold Drift underlies Dreimanis' Plum Point Interstadial beds, which are dated within the time span of the Farmdalian Substage of the Illinois classification. Subdivisions of the Roxana Silt in Illinois may be equivalent to Dreimanis' three units (fig. 1), but, because of the lack of glacial till equivalent to the Dunwich and the lack of finite radiocarbon dates from Illinois older than the time span represented by the Southwold Drift of Ontario, the term Altonian Substage is retained for the entire Roxana sequence.

Thick, calcareous, fossiliferous loess occurs in the valley walls of the Mississippi and Illinois Valleys, but away from these valleys the loess becomes progressively thinner and finer grained (Smith, 1942), is leached, and is nonfossiliferous. Generally the loess deposits are significantly thicker along the east

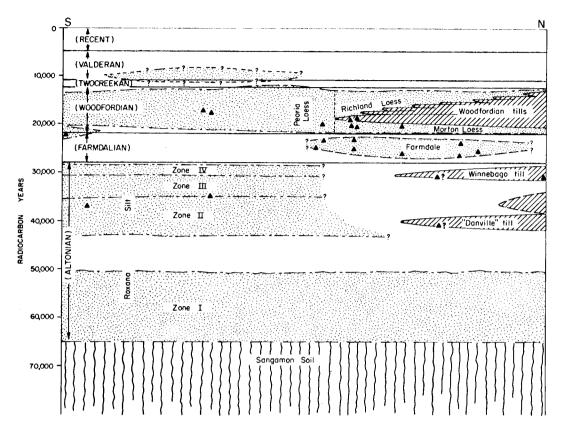


Fig. 1 — Time-space diagram of Wisconsinan deposits in Illinois. The time control by Carbon-14 dates is shown by solid triangles. All radiocarbon dates used were determined in the Washington laboratory of the U.S. Geological Survey. The diagram is arranged north-south, but it is a composite of an east-west belt the entire width of Illinois.

side of the valleys. The asymmetrical distribution of loess with respect to major valleys and the progressive decrease in thickness and grain size away from the valley bluffs generally is also true to the west of Illinois and into the Great Plains region (Kay and Graham, 1943; Frye and Leonard, 1951; Swineford and Frye, 1951) where the thickest loess is related to the valleys of the Missouri and Platte Rivers. Because of this persistent distribution pattern the stratigraphy of the loesses is based largely on the exposures in the major valley walls and the units established there are correlated into the regions of thin, leached loess.

The time and space relations of the loesses in Illinois to the substages of the Wisconsinan are shown by the schematic diagram in figure 1.

Silt Deposits of Altonian Age

The Roxana Silt was named and described by Frye and Willman (1960) from exposures in the east bluff of the Mississippi Valley below the mouths of the Illinois and Missouri Rivers, and includes all of the loess deposits of Altonian age. In its type region the Roxana Silt has been called Late Sangamon loess by Smith

(1942) and Farmdale Loess by Leighton and Willman (1950), but in 1960 Leighton (p. 536) correlated the Peoria Loess at the Roxana type section with the Bignell Loess, Zones II, III, and IV of the Roxana with the Peoria Loess, and Zone I of the Roxana with the Farmdale Loess. However, the radiocarbon date of $35,200\pm1,000$ (W-729) indicates that Zones II and III of the type Roxana are approximately 10,000 years older than type Farmdale, and the weathered Zone I at the base must be still older.

In the Illinois Valley south from Beardstown and along the Mississippi Valley below the mouth of the Illinois, the Roxana is thick and relatively well exposed. In this area distinct lithologic zones are present and can be traced for several hundred miles in a north-south direction. However, as the units are recognizable only where the Roxana is relatively thick and have not been traced into the Ohio, Wabash, or Upper Mississippi Valleys, they are not named as formal rock-stratigraphic units but will be described as informal lithologic zones.

Roxana Zone I is the basal unit of the Roxana. It is more variable and complex than any of the other zones and was weathered before Zone II was deposited. Everywhere that it has been examined it is leached and lacking in fossils, and it commonly rests on Sangamon Soil developed on tills as young as Buffalo Hart (youngest Illinoian). At a few places where Zone I consists largely of silt (e.g., Pleasant Grove Section), it has a weakly developed soil at the top. Commonly, at least the lower part of this zone consists of silty colluvium, with some sand and sparse small pebbles, that grades upward into massive silt. In some places it contains a fairly well sorted sand (e.g., Brown's Mound Section). It is generally brown to gray-tan. In places the basal part is gradational with the top of the underlying Sangamon Soil.

Locally Roxana Zone II overlies Zone I with a transition zone, but generally the contact is sharp (e.g., Cottonwood School Section). Zone II consists of massive, coarse loess that is light to dark pink to pink-tan. Where thick it is weakly dolomitic, and the upper part is fossiliferous. It grades into Zone III above, at some places through a transition of interbedded pink and gray-tan loess. Locally, lenses of eolian sand occur in the upper part.

Roxana Zone III is commonly transitional with the units both above and below. It consists of coarse, massive loess that is calcareous and fossiliferous where thick. It is distinguished by its gray to light yellow-tan color. At several places it contains discontinuous or lenticular bodies of gray to light tan sand, and at some places it is distinguished by abundant large limonite tubules.

Roxana Zone IV (uppermost) is massive, coarse, light pink-tan to gray-pink loess. At localities where Zone III is calcareous, Zone IV is at least weakly calcareous and sparsely fossiliferous in the lower part. However, in most places it is leached in the upper part and at the top has a weathered zone that distinctly separates it from the overlying Peoria Loess. No well developed B-zone has been observed, probably because the interval of weathering was relatively short and because all of the thick sections that have been studied are situated along the valley bluffs where truncation of the Roxana would have occurred during the weathering interval before deposition of the Peoria Loess.

Away from the belt along the Mississippi, Illinois, and Sangamon Valleys, the Roxana is thin and it generally is not possible to differentiate any zones within it. In the Upper Mississippi, Ohio, and Wabash Valleys, the Roxana consists of only a pink-tan to pink-brown loess that is locally sandy and generally grades downward into colluvium that rests on the Sangamon Soil. At only a few places

(e.g., New Harmony Section) have any fossil snails been observed in the Roxana in these areas.

Deposits of Farmdalian Age

The name Farmdale was given by Leighton (1948) to the deposits previously called Late Sangamon. As used here the name Farmdale is restricted to the silt and peat deposits extensively exposed in the type area and dated by numerous radiocarbon dates (pl. 1, fig. 1). The type area of the Farmdale is used as the reference area for the Farmdalian Substage in Illinois. Leighton (1960) recently has applied the name Farm Creek to the peat and wood-bearing silts that have been called Farmdale and are overlain by the Morton ("Iowan") Loess. He reapplied the name Farmdale to some of the older silts, included here within the Roxana, and to the pre-Farmdalian glacial interval.

The Farmdale deposits contain bedded and laminated silt, massive silt, sandy silt, organic-rich silt, and silty organic muck or peat. In color the deposits range from black through various shades of brown and gray-brown to purple. Although generally noncalcareous and nonfossiliferous, at a few places they contain fossil snail shells (Leonard and Frye, 1960).

In some sections (e.g., Richland Creek; Danvers) logs with a maximum diameter up to 3 to 4 inches occur within the Farmdale deposits, but at many places its organic content is merely a humus stain. At the Wedron Section (Leonard and Frye, 1960) the Farmdale includes lacustrine silts that are calcareous, fossiliferous, and contain small flakes of wood.

Loess Deposits of Woodfordian Age

The loess of Woodfordian age is divided into three stratigraphic units distinguished on the basis of their relations to the glacial tills. All loess of this age beyond the limits of Woodfordian glaciation is classed as Peoria. Leverett (1898), p. 186) originally proposed the term Peorian to designate the interval of time between the top of the "Iowan" (Morton) Loess and the overlying Shelbyville till. The term Peorian was subsequently reapplied (Leighton, 1931) to all loess above the "Late Sangamon" (Farmdale Silt) beyond the limits of the Shelbyville Moraine. The Peoria (illustrated by many of the measured sections) is generally gray to yellow-tan, massive loess. It is relatively coarse, calcareous, and fossiliferous adjacent to the major valleys, and locally contains zones or lenticular masses of eolian sand. Away from the valleys it becomes thinner, finer textured, and is leached (e.g., New City and Wolf Creek Sections).

In the lower Illinois Valley and in the Mississippi Valley a weakly developed soil occurs a short distance below the top of the Peoria in a few sections (e.g., Cottonwood School, Frederick South, Collinsville). Radiocarbon dates and regional stratigraphy (pl. 1; fig. 1) suggest the possibility that incipient soils developed during Twocreekan time. However, there is as yet no evidence for excluding this minor soil and the uppermost part of the Peoria Loess from the Woodfordian. Although in the field the upper zone of the Peoria is recognizable only where an incipient soil is present, its mineralogy is distinctive from the main body of the Peoria Loess.

Inside the outermost end moraine (Shelbyville) of the Woodfordian tills, loess occurs both above and below the till. The loess below the Shelbyville till, formerly called Iowan, has been named Morton (Frye and Willman, 1960) from

exposures in Tazewell County. The Morton is gray to yellow-tan, dolomitic, sparsely fossiliferous, massive loess. It ranges from a few inches to as much as 7 or 8 feet thick. Generally it is overlain by Shelbyville till but locally tills as young as Bloomington are in contact with it. At several places a layer of moss has been observed at the top of the loess below the overlying till. The Morton Loess commonly rests on Farmdale deposits but in places is directly in contact with Roxana Silt. The Morton Loess is the oldest stratigraphic unit within the Woodfordian Substage and radiocarbon dates (fig. 1) indicate that its age is from about 20,000 to about 22,000 radiocarbon years B.P. It is stratigraphically continuous with the lower part of the Peoria Loess outside the Shelbyville Moraine.

The Richland Loess (Frye and Willman, 1960), formerly called Tazewell Loess, is the loess that overlies the Woodfordian tills. It is stratigraphically continuous with the upper part of the Peoria Loess. The Richland is gray to yellowtan, and the lower part is calcareous and fossiliferous where thick. Although its maximum thickness exceeds 20 feet, it does not in general exceed 10 to 15 feet. The base of this loess is a time-transgressing stratigraphic boundary as it rests upon tills ranging from Shelbyville, where the loess may exceed 18,000 radiocarbon years B.P., to late Woodfordian, where the base of the loess may be less than 15,000 radiocarbon years B.P. If the Richland contains any loess of Valderan age it has not as yet been recognized.

MINERALOGY

The loesses in Illinois have been studied with regard to thickness, texture, and carbonate content (Smith, 1942), but extensive studies of regional variations or similarities in their mineralogy have not been made previously. Reported here are the results of 185 analyses by optical methods of the minerals in the fine sand (.062 - .250 mm) fraction, 558 analyses by X-ray diffraction of clay minerals, dolomite, and calcite, and 32 chemical analyses of the gross sample. The geographic distribution of samples used in this study is shown by figure 2 and listed in table 1, and the results of analyses are presented in tables 2,4, and 6. In tables 3 and 5 average analyses are listed by depositional province. The positions of the major streams during Altonian time are shown by figure 3.

The mineralogical studies reported here have three general objectives: (1) descriptive characterization of the loesses in Illinois, (2) differentiation of the several stratigraphic units by their mineral composition, and (3) acquisition of information concerning the origin of the loesses by relating their mineral compositions to drainage history and to the composition of the outwash and glacial tills that were their sources.

Clay mineral compositions were calculated from X-ray diffraction data (using oriented aggregate techniques) of the less than 2-micron fraction for montmorillonite, illite, chlorite, and kaolinite. Montmorillonite as used here includes all clay materials that expand to about 17Å with ethylene glycol. Therefore, any expansible chlorite or vermiculite is included with the montmorillonite. Chlorite includes all l4Å material that does not expand with ethylene glycol, and thus includes the non-expansible vermiculite-chlorite and vermiculite. Illite and kaolinite are used as generally accepted.

As kaolinite and chlorite are minor constitutents, their sum rarely exceeding 2 parts in 10, they are listed together in the tables. Montmorillonite and illite, however, may be dominant constitutents and values are given for each.

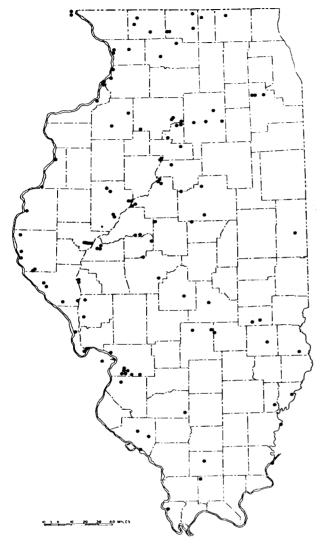


Fig. 2 — Geographic distribution of samples used in this study and listed in tables. Names of measured stratigraphic sections are given on plate 1 and in the described sections.

There is a wide range in the montmorillonite content in the clay mineral assemblages. and there also are differences in relative abundance among the other clay mineral components. In order to characterize these other clay minerals by a numerical value that could be plotted on graphs and tabulated, a diffraction intensity ratio, for convenience called the D.I. ratio. was devised. It was derived by dividing the X-ray diffraction intensity (counts per second) of the 10Å spacing for illite by that for the 7.2A spacing for kaolinite and chlorite. The D.I. ratio has been found to be diagnostic for both calcareous and leached samples of loess, but is not useful in the B-zones of soils or deeply weathered colluvial zones.

Roxana Silt

The mineral composition of the Roxana Silt is shown on plate 1 and in tables 2 and 4. The Roxana generally is low or lacking in calcite, except along the Mississippi Valley below Alton, but Zones II and III and the lower part of Zone IV are dolomitic where the unit is thick. Illite ranges from an average of 14 percent of the clay minerals in the Ancient Mississippi Valley to an average of 38 percent in the Ohio Valley. The D.I. ratios are

consistent, ranging from 0.5 in the Ancient Iowa Valley to 0.8 in the Ohio Valley. Montmorillonite is a major clay mineral in the Roxana but its proportion ranges widely from one depositional province to another. The many samples from the Ancient Mississippi Valley (fig. 3) average 70 percent montmorillonite in the clay minerals, those from the Ancient Iowa Valley average 64 percent, the one section in the Wabash Valley averages 55 percent, and the samples from the Ohio Valley average only 30 percent. For Illinois the amount of illite generally is less than the sum of kaolinite and chlorite.

Quartz is the predominant mineral in the very fine and fine sand fraction in all areas. K-feldspar content ranges from 13 percent to 18 percent, Na-Ca-feldspar

from 7 percent to 10 percent, heavy minerals from less than 0.1 percent to more than 2 percent. The heavy mineral suites along the Ancient Mississippi Valley and the Wabash Valley appear to be similar, with the percentage of hornblende approximately twice that of epidote near the valleys but less in the thin leached sections farther from the valley. In the thin leached sections the relative content of tourmaline, zircon, garnet, and epidote generally is higher.

In the Ancient Iowa Valley staurolite is strikingly more abundant than in any other province, and black opaque minerals, tourmaline, and epidote generally are abundant. The heavy mineral suite in the Ohio Valley is distinctive, being characterized by relatively low content of black opaque minerals, garnet, epidote, and hornblende and relatively high content of tourmaline and zircon, which suggests a mixed source from the bedrock and from glacial outwash.

Farmdale Silt

The Farmdale Silt contains no calcite and no dolomite except at the Wedron Section. At places where the Farmdale lacks organic material, its clay mineral composition is indistinguishable from that of the Roxana

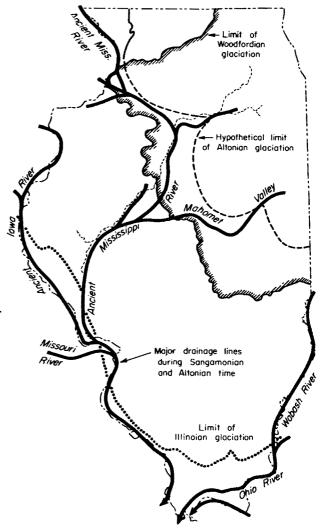


Fig. 3 — Locations and names of major rivers that were sources of loess during Altonian time. Maximum limits are shown for Illinoian and Woodfordian glaciers, and a hypothetical limit of Altonian glaciation also is shown.

of the Ancient Mississippi and Ancient Iowa Valleys (pl. 1, tables 2, 3, 4). Commonly the Farmdale is characterized by an organic-rich appearance and the presence of peaty material, probably indicating local deposition by slow accumulation in a wet environment. It contains montmorillonite, vermiculite, and kaolinite, with little or no illite. Organic-rich Farmdale has an average montmorillonite content of 67 percent in the clay mineral fraction, an average illite content of only 9 percent, and a D.I. ratio of 0.3. The amount of illite is less than the sum of kaolinite and chlorite (vermiculite), similar to that observed for the Roxana. Moreover,

in the Farmdale the amount of vermiculite generally exceeds that of illite. The identical assemblage of clay minerals has been observed in some accretion gleys (Frye, Willman, and Glass, 1960), which suggests similar modes of accumulation and environments of deposition. In a few places the clay minerals of the Farmdale are almost entirely kaolinite and vermiculite. The Farmdale heavy mineral content is 40 percent hornblende, a higher percentage than that of the strongly weathered Roxana Zone I of the same region, but lower than that of the higher zones of the Roxana.

Peoria Loess

The mineral composition of the Peoria Loess permits the recognition of a main body of loess and a discontinuous upper zone. There is greater contrast in mineral composition of the depositional provinces of the Peoria Loess than is found in the Roxana. This contrast is shown strikingly by variations in the amount of montmorillonite in the main body of the loess. Montmorillonite accounts for an average of 70 percent of the clay mineral content in the Mississippi Valley, 54 percent in the Illinois Valley, 26 percent in the Wabash Valley, and a low of 9 percent in the Ohio Valley. In the upper zone of the Peoria Loess an equally strong contrast is evident because montmorillonite composes 80 percent of the clay mineral content in the Mississippi Valley and only 42 percent in the Illinois Valley. In each depositional province the Peoria Loess is more illitic than the Roxana (table 5), and displays a much greater range in illite content - from 11 percent to 71 percent of the clay mineral content. The D.I. ratios of the main body of the loess range from an average 1.2 in the Mississippi Valley above Alton to 3.1 in the Wabash Valley. The average illite content of the Richland Loess and the upper zone of the Peoria of the Illinois Valley ranges from 11 percent of the total clay mineral content in the Mississippi Valley above Alton (D.I. ratio of 0.7) to 41 percent in the Illinois Valley (ratio of 3.0). However, in all areas studied the amount of illite is always greater than the sum of kaolinite and chlorite.

In the very fine and fine sand fraction, the average K-feldspar content ranges from 13 to 16 percent, and the average Na-Ca-feldspar content ranges from 6 to 11 percent. Among the heavy mir erals, zircon content ranges from 3 percent (one sample only) in the Wabash Valley and 4 percent in the Richland Loess of the Illinois Valley to 15 percent in the Mississippi Valley below Alton. Garnet content ranges from an average of 6 percent in the Mississippi Valley above Alton to 19 percent in the Ohio Valley, and the amount of epidote ranges from 10 to 13 percent in the Wabash and Ohio Valleys to 26 percent in the Mississippi Valley below Alton.

Stratigraphic Differentiation

The mineral compositions and the D.I. ratios differentiate the several loesses within each depositional province (pl. 1, figs. 3, 5). The percentage of montmorillonite among the clay minerals of the less than 2-micron fraction is primarily characteristic of the source of sediment in the depositional province and identifies the loesses only within a particular province; their D.I. ratios are reliable indexes to their identity. The average D.I. ratios for the Roxana in all provinces are 0.8 to 0.5, whereas the average ratios for the Peoria in all provinces are 1.2 or higher, except for the upper zone in the Mississippi Valley above Alton where the average is 0.7.

The Farmdale Silt generally has the lowest D.I. ratio of all stratigraphic units studied, but individual samples fall within the range of the Roxana. Where the

Farmdale occurs in an organic-rich environment it has D.I. ratios of 0.2 to 0.3 and the amount of vermiculite exceeds that of illite; both characteristics serve to differentiate it from the Roxana and Peoria. Predominantly vermiculite and kaolinite compositions occur only in the Farmdale.

A transition zone locally occurs in the base of the Peoria (pl. 1) of the Ohio and Wabash Valleys. Its D.I. ratio (1.4) is within the general range of the Peoria Loess of Illinois, but both its D.I. ratio and illite content are significantly lower than those of the Peoria of the Wabash and Ohio Valleys.

Although the clay mineral composition of the upper zone of the Peoria varies widely, the zone, nevertheless, is distinguishable from the main body of the underlying Peoria in the Illinois and Mississippi Valleys by a sharp increase in the D.I. ratio. Where this break is marked by an immature soil (e.g., Collinsville Section on pl. 1) it occurs at the A-zone. In the Illinois Valley and in the Mississippi Valley below Alton the upper zone contains more illite and has a higher D. I. ratio than the underlying Peoria, but in the Mississippi Valley above Alton it contains less illite and more montmorillonite than does the main body of the Peoria.

In thick sections the Roxana Silt generally contains dolomite but only locally contains calcite, whereas calcite is commonly associated with dolomite in the Peoria Loess. The Roxana also generally contains more sand than does the Peoria.

Zone I of the Roxana, which is weathered, has significantly less hornblende and more garnet than do the higher zones. In general, garnet decreases in abundance upward through the Roxana.

Differentiation of Roxana from Peoria by heavy minerals is possible in some areas but not in others, depending on whether the deposits reflect a change in the gross drainage pattern and therefore the ultimate source areas after the deposition of the Roxana. Contrast in heavy mineral suites is most striking in the Mississippi Valley between Alton and Muscatine, where the average percentage of staurolite in the heavy minerals in the Roxana is 8 times that in the Peoria, the percentage of tourmaline is 3 times greater, the percentage of garnet is twice as great, and the percentage of hornblende is only about half as great.

A strong contrast in the heavy mineral suites in the Roxana and Peoria also occurs in the Ohio Valley, where the average percentage of tourmaline in the Roxana is 3 times, zircon 5 times, and garnet and hornblende only half that in the Peoria. The percentage of epidote also is slightly higher in the Roxana. Single samples also suggest a contrast of heavy mineral suites in the Roxana and Peoria in the Wabash Valley. In the Illinois Valley and in the Mississippi Valley below Alton, however, the differences in the heavy mineral suites are not sufficiently marked or consistent to serve as criteria for the differentiation of stratigraphic units.

To summarize, wherever the Roxana, Farmdale, and Peoria deposits are not strongly weathered, their total mineral assemblages provide reliable criteria for their differentiation, for the distinction of Zone I of the Roxana from the higher zones, and for the subdivision of the Peoria Loess.

ORIGIN

It is not our purpose here to review the long controversy relative to the eolian versus non-eolian origin of loess, as in our judgment the loess of the Midwest has been demonstrated to have an eolian origin (e.g., Udden, 1894, 1898; Smith, 1942; Leighton and Willman, 1950; Kay and Graham, 1943; Swineford and Frye, 1951, 1955; Leonard and Frye, 1954). We are concerned here with the source of the

materials that compose the various loess sheets of Illinois.

The immediate source of loess is the fine-size fraction of the water-transported and water-deposited sediments of major valleys. The bulk of these sediments, however, was produced as outwash from continental glaciers, and to understand properly the ultimate source of the loess we must know the mineral composition of the glacial tills related to the outwash. However, the till related to the outwash is not the till now occurring immediately below the loess. Rather, it is the till that occurs up the drainage way, possibly a considerable distance from the loess deposit. It is therefore requisite that the courses of major streams carrying outwash during the several episodes of loess deposition be known.

Drainage History

Changes in the drainage patterns of the Upper Mississippi Basin during early Pleistocene time were complex and their history is controversial, but here we are concerned only with the patterns existing during post-Illinoian time (fig. 3). The major stream in Illinois was the Ancient Mississippi River, which coincided with the present Mississippi past Jo Daviess and Carroll Counties but in Whiteside County flowed southeastward to the position of the present valley of the Illinois River in southeastern Bureau County. From there the river followed the general trend of the Illinois southward. It is shown in figure 3 as parallel to, but 5 to 15 miles east of, the present Illinois River in parts of Woodford, Tazewell, and Mason Counties, returning to the present Illinois Valley by way of the lower Sangamon Valley, which was continuous with the Mahomet River from the east. It has been tentatively concluded by George E. Ekblaw (personal communication) that the present narrow segment of the valley at Peoria was initiated by the Jacksonville Stage of the Illinoian glaciation and was later overridden, but not deeply buried, by the Buffalo Hart till. It may be that both the eastern and western valleys were relatively open until the advance of the Shelbyville (earliest Woodfordian) glacier. Southward from Cass County the Ancient Mississippi followed the present Illinois Valley and then the present Mississippi Valley to the north end of what later became the Thebes Gorge.

The Ancient Iowa River, which drained the basin of the present Iowa River in Iowa, followed the present Mississippi Valley from Muscatine southward, gaining tributaries from the basin of the Des Moines River in Iowa and joining the Missouri River and the Ancient Mississippi at the mouth of the present Illinois River.

The Ohio River occupied its present valley to below Golconda and then flowed westward across northern Massac and Pulaski Counties by way of the Cache Valley. The Wabash occupied the existing valley along the eastern boundary of Illinois.

This drainage pattern is judged to have remained essentially unchanged from the time of the withdrawal of the Illinoian glaciers, throughout Sangamonian, Altonian, and Farmdalian time, until it was disrupted by the advancing Woodfordian glaciers. As shown in figure 3, the Altonian glaciers of the Lake Michigan lobe approached but did not cross the Ancient Mississippi River, although the data are far from conclusive. Thus, during Altonian time (when Roxana Silt was deposited) the Ancient Mississippi River could have carried outwash as well as locally derived alluvium from Minnesota and northern Iowa while also receiving local alluvium from Illinois and outwash from the Altonian glaciers that entered Illinois from the Lake Michigan lobe and perhaps from as far east as the Saginaw lobe.

In contrast, after the major stream was blocked in Bureau and Whiteside Counties by advancing Woodfordian glaciers (Shaffer, 1954), the part of this valley that later became the lower part of the Illinois River Valley received outwash only from the Lake Michigan lobe. This point in time occurs within the Morton Loess. Furthermore, the segment of the Mississippi Valley between Alton and Muscatine received sediments from only central and southeastern Iowa and northeastern Missouri during Altonian and Farmdalian time, but after the major diversion during Woodfordian time it received sediments from the entire upper part of the present basin of the Mississippi River.

The Wabash Basin received outwash during Altonian time from the glacier that deposited the till called "Farmdale" by Ekblaw and Willman (1955) at Danville, Illinois, and perhaps also from lobate fronts farther east in Indiana. In Woodfordian time the Wabash Basin received contributions largely from the glacial lobes that advanced southwestward across Indiana from Michigan and Ontario. Throughout Wisconsinan time the Ohio Valley east of the Wabash Valley received no sediment from a source in the Lake Michigan lobe or farther west.

Throughout Wisconsinan time the Missouri River also was contributing outwash to the Mississippi Valley below Alton from glaciers as far west as the Rocky Mountains.

Mineralogy of Glacial Tills

The mineralogy of the tills of Illinois is discussed in the forthcoming companion Circular to this report, but some generalities are summarized here. In many areas some of the older till was incorporated in the overriding younger glaciers, and beyond the limits of the younger glaciers it furnished sediments to the valleys that carried outwash from local erosion.

In general, the clay fraction of the tills west of Illinois is dominantly montmorillonite with minor amounts of illite and kaolinite, whereas east of Illinois, the clay fraction of the tills is dominantly illite and chlorite with minor amounts of montmorillonite. In northeastern Kansas (Iowa Point Section) the clay minerals of Nebraskan and Kansan tills contain more than 80 percent montmorillonite, whereas those in Kansan till of western Indiana (Cagle's Mill Section) contain less than 10 percent montmorillonite. In Illinois, the clay minerals of the western-source Kansan till west of the Illinois Valley are 60 to 70 percent montmorillonite.

The Illinoian tills in Illinois all contain some kaolinite and chlorite, whereas the percentages of illite and montmorillonite range widely. The clay mineral composition of Illinoian tills of extreme western Illinois approaches (45 to 65 percent montmorillonite) that of the Kansan till. In general the Jacksonville and Payson tills in the Illinois Valley region have a clay mineral fraction that is 20 to 30 percent montmorillonite. These tills in eastern Illinois contain about 15 percent montmorillonite, whereas the younger Illinoian Buffalo Hart till contains only 0 to 6 percent montmorillonite.

The clay mineral composition of the Wisconsinan tills of Illinois ranges less widely. The clay minerals in the mid-Altonian till at Danville, which presumably furnished outwash to the Wabash Valley, consist of 26 percent montmorillonite, 50 percent illite, and 24 percent kaolinite and chlorite. The late Altonian Winnebago till of north-central and northwestern Illinois has a clay mineral content that is up to 30 percent montmorillonite. It also contains prominent illite, some chlorite, but no kaolinite. The montmorillonite content in the clay minerals of older Woodfordian tills (Shelbyville-Bloomington) ranges from 0 to 15 percent. Illite and chlorite are the dominant clay minerals. In the eastern part of the state montmor-

illonite is present only in the outer part of the Shelbyville Moraine. Illite and chlorite predominate in Woodfordian tills younger than Bloomington; kaolinite and montmorillonite generally are absent.

Glaciers, therefore, that invaded Illinois from the northwest or those that invaded from the northeast and overrode tills that had been deposited by glaciers from the northwest furnished outwash sediments high in montmorillonite; glaciers entering Illinois from the northeast furnished outwash sediments high in illite and chlorite.

Mineralogy of Glacial Outwash

Outwash of Wisconsinan age was sampled in the Illinois and Mississippi Valleys. These samples range in texture from coarse sand and gravel to silty clay, and most of them yielded sufficient clay minerals for X-ray analysis. The montmorillonite content of the clay minerals segregated from Altonian outwash in the Mississippi Valley above Alton ranges from 29 to 72 percent and averages 56 percent, that from Woodfordian outwash ranges from 47 percent to 81 percent and averages 68 percent, and that from Valderan outwash averages 71 percent. In the clay minerals from 30 samples of Woodfordian outwash from the Illinois Valley, montmorillonite content ranged from 0 (six samples) to 60 percent, with an average of 24 percent. Kaolinite and chlorite generally constitute from 10 to 20 percent of the clay minerals present, and illite is the predominant clay mineral in Woodfordian outwash of the Illinois Valley. In the heavy mineral assemblages (tables 2, 3) outwash from the Mississippi Valley is distinctly higher in epidote than that from the Illinois Valley. Garnet is relatively high in Illinois Valley outwash.

Roxana Silt

With this general regional background, the sources of the materials in the several stratigraphic units within the loesses of the depositional areas is considered. The sequence of sections along the Ancient Mississippi River from DePue to Gale, plate 1, shows the generally high montmorillonite content of the Roxana. At the north end of the traverse the high montmorillonite content shows that the chief source of sediment was from the northwest. From the Richland Creek through the Farm Creek Sections the Roxana is thin and weathered. The lowest montmorillonite percentage in the Roxana along the Ancient Mississippi Valley occurs at Frederick South where the Roxana is quite sandy. The montmorillonite content irregularly increases southward because of the increasing contribution of loess from the Ancient Iowa Valley that converged from the northwest, and reaches a maximum (French Village Section) below the mouth of the Missouri River which contributed sediments from the montmorillonite-rich western areas. Calcite is more common in the Roxana Silt south of the Missouri River, because the western tills are high in limestone whereas those from the Lake Michigan lobe are high in dolomite.

The other sequence of sections on plate 1 crosses the major valleys and therefore illustrates the differences in composition of Roxana Silt derived from different source areas. At the north, the Fulton Quarry Section with its high percentage of montmorillonite is in the Ancient Mississippi Valley and reflects a northwestern source. The next three sections are in the Ancient Iowa Valley which drained a different and more local source area that provided sediments with a slightly lower montmorillonite content. At Frederick South and New City the composition of the Roxana is again controlled by the Ancient Mississippi Valley source. At New Harmony

in the Wabash Valley the Roxana Silt has a slightly lower montmorillonite content. At the Newburg and Yankeetown Sections of the Ohio Valley the amount of montmorillonite in the Roxana Silt reaches its lowest value, whereas in The Rocks Section, in the Ohio Valley just below the mouth of the Wabash, it is intermediate to that in the deposits of the two valleys.

It seems probable that Zones II, III, and IV of the thick Roxana along the Ancient Mississippi River were deposited during successive advance and retreat of Altonian glacial lobes (fig. 1). The essentially similar mineralogy of the three zones indicates that the successive glaciers occupied similar areas and moved in identical directions. In contrast, Zone I generally differs significantly in mineralogy from the higher zones and generally is weathered, in some places intensely so. Below the mouth of the Missouri River, Roxana Zone I commonly contains leached loess gradationally above colluvium at the base, but regionally Zone I generally is sandy colluvium with little loess. These relations lead to the tentative conclusion that Zone I was produced largely from relatively local sources as the result of climatic change, and that if glaciers of this age invaded the basins of the Ancient Mississippi and Ancient Iowa Rivers they penetrated only short distances.

X-ray determinations of calcite and dolomite were made both on the less than 2-micron fraction and on the bulk sample (table 4). Dolomite was detected in the bulk sample taken at some localities from the top few inches of the Roxana but was not detected in the less than 2-micron fraction. It seems likely that in such places this thin zone is either the result of mixing during deposition or is the first slowly accumulating deposit of Woodfordian loess, leached of the very fine dolomite but retaining some of the silt-size dolomite grains.

At a few places (e.g., New Harmony Section) snail shells are sparsely present in loess that does not otherwise contain carbonate minerals. The effect produced by differences in grain size on the leaching of carbonate minerals was checked experimentally in the laboratory and it was demonstrated that not only could all the silt-size calcite be removed before the dolomite was appreciably affected, but the silt-size dolomite could be removed before massive snail shells were noticeably affected. These solubility differences explain the virtual absence of calcite in the fine fractions of the Roxana Silt.

Farmdale Silt

The Farmdale Silt shown on plate 1 occurs only in the Ancient Mississippi Valley. Generally, wherever it is thick and rich in organic matter it lies on thin, weathered Roxana Silt (pl. 1), indicating that the Farmdale Silt accumulated slowly in relatively poorly drained situations that did not exist in areas of thick Roxana. Such conditions probably account for the distinctive mineralogy of the Farmdale Silt, as minerals were supplied largely by local sheet wash from Roxana Silt and older tills, probably supplemented by a minor amount of eolian silt from the Ancient Mississippi Valley. The fact that this increment of loess was not large and was deposited slowly is consistent with its accumulation when the glaciers were remote and little if any outwash was being transported down the major valley, and with the fact that the Farmdale silts generally were laminated. A small amount of Farmdale loess with no organic matter may be indistinguishably included in the top of the Roxana Silt.

Loesses of Woodfordian Age

Loesses of Woodfordian age in Illinois include the Morton Loess, the Richland Loess, the Peoria Loess (except perhaps the uppermost zone, pl. 1, fig. 1), and at least part of the transition zone at the base of the Peoria in the Ohio Valley.

During the initial pulse of loess deposition in Woodfordian time, the Ancient Mississippi River (fig. 3) was carrying sediment from as far north as Minnesota. The region that is now the Illinois Valley was deprived of this northwestern source when the advancing Shelbyville glacier blocked the river in Bureau County and forced the drainage from the north to flow across the low divide above Muscatine. This diverted the uppermost part of the Ancient Mississippi River into the Ancient Iowa Valley and established the present course of the Mississippi River. The Woodfordian loess of the Illinois Valley strikingly records this change in sediment source that deprived the valley of the large contribution of montmorillonite from the northwest and restricted the outwash to the high illite-chlorite contribution that came largely from the advancing Lake Michigan lobe.

The Morton Loess, deposited in the valley during glacial advance and overlain by Shelbyville till, as well as the basal part of the Peoria Loess beyond the Shelbyville Moraine, also reflects this history. The upper boundary of the Morton Loess is time-transgressing because loess deposition continued during the advance of the glacial front from DePue to the outermost limit of the Shelbyville Moraine. At DePue a high montmorillonite content throughout the Morton reflects only the Ancient Mississippi Valley source. In the Farm Creek Railroad Cut Section the clay minerals in the lower part of the Morton, which is derived from the Ancient Mississippi Valley, contain 70 percent montmorillonite and only 17 percent illite, but the clay minerals in the upper part of the Morton, which was derived from the Illinois Valley after the diversion, contain only 29 percent montmorillonite and 46 percent illite. The low montmorillonite content of the Morton in the Illinois Valley south of the DePue Section reflects the blocking of the Ancient Mississippi River, as most of the Morton samples from this valley segment are from the upper part of the loess. The fact that the montmorillonite content in the lowermost samples from the Morton and Peoria Loesses of the Illinois Valley is the same as the average montmorillonite content of the main part of the Peoria in the Mississippi Valley suggests that these loesses had a similar source.

The main body of the Peoria Loess along the Illinois Valley contains significantly less montmorillonite (54 percent) than does that along the Mississippi Valley, but the much lower montmorillonite content (pl. 1) in the Peoria in the Wabash Valley (26 percent) and in the Ohio Valley (9 percent) reflects the paucity of this mineral in the glacial tills and bedrock of the Ohio River basin east of the Wabash Valley. Also, the fact that the mineral composition at The Rocks Section in the Ohio Valley, just below the mouth of the Wabash, is virtually an average of the loesses of these two valleys above their junction, demonstrates that the loesses, including their clay minerals, are derived from the adjacent source valleys.

The abundance of montmorillonite in the loesses in the Mississippi and Illinois Valleys is therefore a consequence of their derivation from the sediments in the valleys and can be accounted for without invoking long distance eolian transport from remote western sources (Beavers et al. 1955; Beavers, 1957).

The mineral composition of the Peoria Loess is consistent across the broad upland divide areas of south-central Illinois and is closely similar to the composition of the loess in the bluffs of the Illinois or Mississippi Valleys to the west-

northwest. The decrease in grain size of the loess toward the east-southeast is accompanied by a slight increase in the percentage of montmorillonite in the clay mineral fraction. This slight increase may be accounted for by the finer particle size of the montmorillonite among the clay minerals of the source valleys. The admixture of sediments rich in illite and chlorite from the Wabash Valley, the first major source area to the east, is not recognizable in the Peoria Loess more than 20 to 30 miles to the west of the Wabash Valley.

The transition zone in the Ohio Valley appears to be derived from outwash from the advancing Woodfordian glacier. The Erie lobe probably would have entered the Ohio Valley somewhat earlier than the Michigan or Superior lobes entered the Illinois or Ancient Mississippi Valleys. The radiocarbon date supports this concept. The mineralogy indicates that the source of the transition zone must have also included materials derived from the Roxana and older deposits of the valley.

Although far less striking in their contrasts, the heavy mineral assemblages (table 3) of the Woodfordian loesses also reflect the source areas.

CONCLUSIONS

The position of Illinois on the North American continent is well suited for demonstration of the concept that the minerals in the loess were derived from the outwash transported through the major valleys. Illinois received sediments both from the regions to the west and northwest where montmorillonite predominates among the clay minerals and from the regions to the east and northeast where illite and chlorite predominate. Furthermore, drainage modifications during the Wisconsinan age resulted in shifting the sediment sources of central Illinois loess between these two regions.

From the data presented in this report, several general conclusions can be drawn.

- (1) The source of the outwash from which the loess was in part derived may be hundreds of miles up the major valleys from the point of loess deposition, although the rocks adjacent to these valleys may furnish enough sediment by local erosion to produce a detectable effect on the relative abundance of mineral species in the loess.
- (2) An abrupt change in the source of sediments, caused by the blocking of a valley carrying outwash, is sharply reflected in the mineral composition of the down-valley loess if the source areas involved possessed significantly different mineral assemblages.
- (3) The non-carbonate mineral composition of a loess below the B-zone of the surface soil persists as a recognizable assemblage across the region east and southeast of one major source valley to the proximity of the next source valley.
- (4) Both the clay and the coarser fractions of the loess were derived locally from the sediment transported through the source valleys.
- (5) Differential leaching of calcite and dolomite, and differential leaching of very fine dolomite and silt size dolomite, both in transit and in situ, produced marked differences in the carbonate content of loess.
- (6) The differences in mineral composition of the several stratigraphic units in the loesses of Illinois can be used for identification of the units.

REFERENCES

- Beavers, A. H., 1957, Source and deposition of clay minerals in Peoria loess: Science, v. 126, p. 1285.
- Beavers, A. H., Johns, W. D., Grim, R. E., and Odell, R. T., 1955, Clay minerals in some Illinois soils developed from loess and till under grass vegetation: Clays and Clay Minerals, Proceedings of the Third Natl. Conf. on Clays and Clay Minerals; Pub. 393, Natl. Acad. of Sci. Natl. Research Council, Washington, D. C., p. 356-372.
- Chamberlin, T. C., 1897, A supplementary hypothesis respecting the origin of the loess of the Mississippi Valley: Jour. Geology, v. 5, p. 795-802.
- deVries, Hessel, and Dreimanis, Aleksis, 1960, Finite radiocarbon dates of the Port Talbot interstadial deposits in southern Ontario: Science, v. 131, no. 3415, p. 1738-1739.
- Dreimanis, Aleksis, 1960, Pre-classical Wisconsin in the eastern portion of the Great Lakes region, North America: Internatl. Geol. Cong., 21st Session, Norden, pt. IV, p. 108-119.
- Dreimanis, Aleksis, 1961, Supplement to Pre-classical Wisconsin in the eastern portion of the Great Lakes region, North America: Univ. Western Ontario, Dept. Geol., Contr. 33a (mimeo.), 6 p.
- Ekblaw, G. E., and Willman, H. B., 1955, Farmdale drift near Danville, Illinois: Illinois Acad. Sci. Trans., v. 47, p. 129-138.
- Frye, J. C., and Leonard, A. B., 1951, Stratigraphy of the late Pleistocene loesses of Kansas: Jour. Geology, v. 59, p. 287-305.
- Frye, J. C., and Willman, H. B., 1960, Classification of the Wisconsinan Stage in the Lake Michigan glacial lobe: Illinois Geol. Survey Circ. 285, p. 16.
- Frye, J. C., Willman, H. B., and Glass, H. D., 1960, Gumbotil, accretion-gley, and the weathering profile: Illinois Geol. Survey Circ. 295, p. 39.
- Horberg, Leland, 1953, Pleistocene deposits below the Wisconsin drift in northeastern Illinois: Illinois Geol. Survey Rept. Inv. 165, p. 61.
- Horberg, Leland, 1956, Pleistocene deposits along the Mississippi Valley in central-western Illinois: Illinois Geol. Survey Rept. Inv. 192, p. 39.
- Kay, G. F., and Graham, J. B., 1943, The Illinoian and post-Illinoian Pleistocene geology of Iowa: Iowa Geol. Survey, v. 38, p. 1-262.
- Leighton, M. M., 1917, The Iowan glaciation and the so-called Iowan loess deposits: Iowa Acad. Sci. Proc., v. 24, p. 87-92.
- Leighton, M. M., 1931, The Peorian loess and the classification of the glacial drift sheets of the Mississippi Valley: Jour. Geology, v. 39, p. 45-53.
- Leighton, M. M., 1948, Footnote on p. 390, in Wascher, H. L., Humbert, R. P., and Cady, J. G., Loess in the southern Mississippi Valley: Identification and distribution of the loess sheets: Soil Sci. Soc. Amer. Proc. 1947, v. 12, p. 389-399.

- Leighton, M. M., 1960, The classification of the Wisconsin Glacial Stage of North Central United States: Jour. Geology, v. 68, p. 529-552.
- Leighton, M. M., and Willman, H. B., 1950, Loess formations of the Mississippi Valley: Jour. Geology, v. 58, p. 599-623.
- Leonard, A. B., and Frye, J. C., 1960, Wisconsinan molluscan faunas of the Illinois Valley region: Illinois Geol. Survey Circ. 304, p. 32.
- Leonard, A. B., and Frye, J. C., 1954, Ecological conditions accompanying loess deposition in the Great Plains region of the United States: Jour. Geology, v. 62, p. 399-404.
- Leverett, Frank, 1898, The weathered zone (Sangamon) between the Iowan loess and the Illinoian till sheet: Jour. Geology, v. 6, p. 171-181.
- Leverett, Frank, 1899, The Illinois glacial lobe: U.S. Geol. Survey Monograph 38, p. 1-817.
- Savage, T. E., 1915, The loess in Illinois, its origin and age: Illinois Acad. Sci. Trans., v. 8, p. 100-117.
- Shaffer, P. R., 1954, Extension of Tazewell Glacial Substage of western Illinois and eastern Iowa: Geol. Soc. Amer. Bull., v. 65, p. 443-456.
- Smith, G. D., 1942, Illinois loess Variations in its properties and distribution: A pedologic interpretation: Univ. of Illinois Agr. Experiment Sta. Bull. 490, p. 139-184.
- Swineford, Ada, and Frye, J. C., 1951, Petrography of the Peoria loess in Kansas: Jour. Geology, v. 59, p. 306-322.
- Swineford, Ada, and Frye, J. C., 1955, Petrographic comparison of some loess samples from western Europe with Kansas loess: Jour. Sed. Petrology, v. 25, p. 3-23.
- Udden, J. A., 1894, Erosion, transportation, and sedimentation performed by the atmosphere: Jour. Geology, v. 2, p. 318-331.
- Udden, J. A., 1898, The mechanical composition of wind deposits: Augustana Libr. Pub. 1, 69 p.

Sample no.	Location 1/4 Sec., T., R., County	Stratigraphic unit	Thickness of unit	Feet below top of unit
8	Center SE, 20-3N-8W, Madison	Roxana-II	20.0	16.0
8-A	do.	Roxana-I	4.0	3.5
8-AA	do.	do.	4.0	3.0
8-B	do.	do.	4.0	2.0
8-BB	do.	do.	4.0	1.0
10	do.	Roxana-II	20.0	5.0
10-A	do.	Roxana-III	12.0	7.0
11	do.	Peoria	11.0	6.0
12	do.	Roxana-IV	15.0	10.0
17	SW SE NE, 31-3N-6W, Madison	Roxana	7.5	5.0
18	do.	do.	7.5	2.5
19	do.	Peoria	5.0	3.0
20	SE SW NE, 10-5N-10W, Madison	Peoria	12.0	9.0
21	do.	Roxana	19.0	10.0
22	do.	do.	19.0	16.5
23	do.	Roxana-I	2.0	1.0
30	Ft. Bellfontaine, St. Louis Co., Missouri	Roxana	10.0	8.0
31	do.	do.	10.0	5.5
32	do.	do.	10.0	3.0
33	do.	do.	10.0	1.0
34	do.	Peoria	15.0	14.5
35	do.	do.	15.0	8.0
36	do.	do.	15.0	4.0
39	SE SE, 32-10S-2W, Calhoun	Roxana	2.5	1.5
40	do.	Peoria	15.0	11.0
83	Center SE NW, 13-3N-8W, Madison	Peoria	13.0	9.0
84	do.	Roxana	14.0	5.0
85	do.	do.	14.0	9.5
86	do.	do.	14.0	12.0
87	NW NE SW, 32-3N-7W, Madison	Peoria	10.0	7.0
88	do.	Roxana	5.5	1.5
89	do.	do.	5.5	3.5
99	Center E. line, 11-18N-11W, Cass	Roxana	38.0	36.0
100	do.	do.	38.0	31.0
101	do.	do.	38.0	27.0
102	do.	do.	38.0	23.0
103	do.	do.	38.0	17.0
104	do.	do.	38.0	13.0
105	do.	do.	38.0	9.0
106	do.	do.	38.0	2.0
107	do.	Peoria	40.0	38.0
108	do.	do.	40.0	32.0
109	do.	do.	40.0	27.0
110	do.	do.	40.0	20.0
115	NE SE NE, 10-6N-5E, Fulton	Roxana	7.0	6.0
116	do.	do.	7.0	4.0
117	do.	do.	7.0	1.0
118	do.	Peoria	18.0	16.0
127	NE NW NE, 26-26N-4W, Tazewell	Farmdale	6.0	4.0
128	do.	do.	6.0	1.0

TABLE 1 - Continued

Sample no.	Location 1/4 Sec., T., R., County	Stratigraphic unit	Thickness of unit	Feet below top of unit
129	NE NW NE, 26-26N-4W, Tazewell	Morton	2.5	1.0
133	Center, 31-26N-3W, Tazewell	Roxana	0.8	0.4
134	do.	Farmdale	3.5	2.5
135	do.	do.	3.5	0.5
136	do.	Morton	4.0	3.0
137	do.	do.	4.0	1.0
140	SW NW SW, 12-26N-4W, Tazewell	Richland	10.0	9.0
144	SW NW NE, 1-25N-2E, McLean	Farmdale	2.5	1.0
152	SW NW SE, 15-36N-9E, Will	Valparaiso ow.*		***
153	NE NE NW, 16-34N-9E, Will	do.		
154	NE NW SE, 32-14N-10E, Putnam	Woodf. ow.		
155	NE NW SW, 17-11N-9E, Peoria	Woodf. ow.		
156	SE SW SW, 16-23N-7W, Tazewell	do.		
157	do.	dune sand		
158	SW SW SW, 20-18N-11W, Cass	do.		
159		Valparaiso ow.		
160	SE SE SE, 8-46N-2E, Winnebago NE NW NW, 5-44N-2E, Winnebago	do.		
161	NW SE NE, 29-25N-11E, Ogle	Winnebago ow.		
162	NW NE NW, 27-17N-5E, Henry	Dune sand		
163	SW SE NE, 23-24N-3E, Carroll	Mankato ow.		
164	SE NE SW, 34-21N-2E, Rock Island	Mankato ow.		
165	NW SW, 4-20N-2E, Rock Island	Dune sand		
166	NE NE SW, 15-11N-5W, Henderson	Mankato ow.		
175	NW SW NW, 30-33N-1W, Putnam	Woodf. ow.		
176	SE NE SE, 21-37N-12E, Cook	Lemont ow.		
177	NW SE NW, 36-30N-3W, Marshall	Woodf. ow.		
178	NE NW SW, 26-33N-1W, Putnam	do.		
179	NW SW NW, 30-33N-1W, Putnam	do.		
271	NE NW SE, 16-7S-2W, Pike	Roxana	4.0	3.0
272	do.	Peoria	8.0	6.5
281	NE SE, 27-14N-13W, Scott	Roxana	21.0	4.0
282	do.	Peoria	25.0	21.0
283	do.	do.	25.0	10.0
284	do.	do.	25.0	25.0
287	SW SE NW, 26-14N-13W, Scott	Peoria	25.0	15.0
288	do.	do.	25.0	8.0
310	NE NE NW, 10-18N-1E, Rock Island	Roxana	3.0	1.5
311	do.	Peoria	10.0	5.0
319	SW NW NW, 15-15N-5E, Henry	Peoria	10.0	6.0
324	SE cor. NE NW, 35-15N-6E, Bureau		3.0	1.5
325	do.	Peoria	15.0	14.0
326	do.	do.	15.0	7.0
332	SE NE NW, 36-8N-2E, Fulton	Roxana	0.8	0.4
333	do.	Peoria	10.0	7.0
341	SE NE SE, 4-14N-4W, Sangamon	Roxana	3.0	1.5
342	do.	Peoria	6.0	4.5
342-A	do.	Roxana	3.0	2.5
342-B	do.	do.	3.0	2.0
342-C	do.	do.	3.0	1.0
342-D	do.	do.	3.0	0.5
J42 D	uo.	ΨO.	3.∪	0.5

TABLE 1 — Continued

Sample no.	Location 1/4 Sec., T., R., County	Stratigraphic unit	Thickness of unit	Feet below top of unit
342-E	SE NE SE, 4-14N-4W, Sangamo	n Peoria	6.0	5.5
342-F	do.	do.	6.0	5.0
342-G	do.	do.	6.0	4.0
342-H	do.	do.	6.0	3.5
342-I	do.	do.	6.0	2.7
342-J	do.	do.	6.0	1.5
351	NW SE NE, 29-22N-4E, McLean		4.0	3.5
352	do.	do.	4.0	2.5
359	NW SE NW, 6-21N-4W, Mason	Peoria	7.0	6.5
360	do.	do.	7.0	5.0
361	SW SE SW, 2-23N-7W, Tazewel	1 Woodf. ow.	35.0	25.0
362	do.	do.	35.0	15.0
366	NW SE NW, 17-11N-9E, Peoria		70.0	50.0
373	SE SE SW, 15-16N-10E, Burea		5.0	4.0
379	SW SE SE, 27-16N-10E, Burea		1.3	0.5
380	do.	Roxana-II	1.0	0.5
381	do.	Roxana-III	4.0	2.0
382	do.	Roxana-IV	1.2	0.6
383	do.	Morton	2.0	1.0
385	do.	Lake Ill. silt	4.0	2.0
386	SE SW NE, 35-16N-11E, Burea	u Woodf.ow. (silt bed)	1.8	1.0
387	do.	Woodf. ow.	35.0	15.0
389	SW SW SW, 18-33N-3E, LaSalle		15.0	10.0
393	NE NE NW, 22-36N-10E, Will	Valparaiso ow.	30.0	4.0
406	NE NW NE, 28-46N-6E, McHenr	• .	1.0	0.5
414	NE SW NW. 32-46N-3E. Boone	Peoria	1.0	0.9
420	Center E. line, 5-26N-10E, Winnebago	do.	4.0	3.0
425	SW SW NW, 6-26N-8E, Stephen	son do.	6.0	5.0
426	NE NE SE, 5-27N-6E, Stephen	son Roxana	2.0	1.0
427	do.	Farmdale	0.5	0.3
428	do.	Peoria	2.5	2.0
432	SE SW SW, 14-81N-6E, Clinton Iowa	n, do.	40.0	30.0
433	do.	do.	40.0	35.0
434	NW NE NE, 21-21N-3E, Whites		15.0	5.0
435	SE NW SW, 34-21N-2E, Rock I	sland Woodf. ow.	15.0	3.0
436	do.	do.	15.0	5.0
452	SW SW SW, 16-3S-1W, Washing	ton Roxana	5.0	3.0
457	NW NW NE, 32-7S-6W, Randolp	h Peoria	15.0	11.0
460	Center, 33-14S-3W, Alexande	r Roxana-I	6.0	5.0
461	do.	do.	6.0	3.0
462	do.	do.	6.0	1.0
463	do.	Roxana-II	7.0	6.0
464	do.	do.	7.0	1.0
465-A	do.	Roxana-III	8.0	7.5
465-B	do.	do.	8.0	6.0
466	do.	do.	8.0	0.5
467	do.	Roxana-IV	6.0	2.5
468	do.	do.	6.0	0.3

Sample no.	Location 1/4 Sec., T., R., County	Stratigraphic unit	Thickness of unit	Feet below top of unit
470		Peoria	25.0	23.5
	Center, 33-14S-3W, Alexander	do.	25.0 25.0	18.0
471	do.		25.0 25.0	12.0
472	do.	do.		
473	do.	do.	25.0	6.0
477	1 mi. SE Shawneetown bridge, Ky.	Roxana	6.0	3.0
478	do.	do.	6.0	0.5
479	do.	Peoria(trans.)	5.0	2.5
480	do.	Peoria	20.0	18.0
481	SW SE NW, 26-7S-4W, Pike	Peoria	12.0	6.0
483	Center NE, 25-19N-5W, Menard	Roxana	3.0	1.5
484	do-	Peoria	10.0	9.0
485	NW NW NE, 33-20N-6W, Mason	Roxana	3.0	2.5
486	do.	do.	3.0	1.0
487	do.	Peoria	12.0	11.5
488	do.	do.	12.0	7.0
489	SE SE SW, 9-34N-4E, La Salle	Farmdale	35.0	25.0
490	do.	do.	35.0	12.0
491	do.	do.	35.0	9.0
493	do.	do.	35.0	2.0
501	SE NW NE, 31-85N-6E, Jackson, Iowa	Peoria	6.0	5.0
504	SW NW SE, 1-88N-2E, Dubuque, Iowa	Roxana	4.0	2.0
505	do.	Peoria	20.0	18.0
511	Center N. line, NW, 32-89N-2E,	Roxana	3.5	2.5
	Dubuque, Iowa			
512	do.	do.	3.5	1.0
513	do.	Peoria	15.0	14.0
514	do.	do.	15.0	8.0
515	do.	do.	15.0	6.0
523	SW cor. NW, 11-24N-3E, Carroll	Roxana	8.0	7.0
524	do.	do.	8.0	4.0
525	do.	do.	8.0	1.5
526	do.	Peoria	25.0	23.0
527	do.	do.	25.0	8.0
533	NE NE SW, 21-22N-3E, Whiteside	Roxana-I	4.7	2.0
534	do.	do.	4.7	1.0
535	do.	do.	4.7	0.2
	40.			
536	do.	Peoria	35.0	33.0
537	do.	do.	35.0	5.0
544	NW NE SE, 15-21N-1E, DeWitt	Roxana	4.5	4.0
545	do.	do.	4.5	2.5
546	do.	do.	4.5	1.0
547	do.	Morton	3.0	2.5
548	do.	do.	3.0	0.5
553	NW NE NW, 32-25N-1W, Woodford	Roxana	1.5	1.0
554	do.	Farmdale	5.5	4.5
555	do.	do.	5.5	3.0
		-		
556	do.	do.	5.5	1.0
557	do.	Morton	2.6	0.5
561	Center SE NE, 20-28N-2W, Woodford	Roxana	3.0	2.0
562	do.	Farmdale	7.5	6.5
563	do.	do.	7.5	4.0

TABLE 1 — Continued

Sample no.	Location 1/4 Sec., T., R., County	Stratigraphic unit	Thickness of unit	Feet below top of unit
564	Center SE NE, 20-28N-2W, Woodford	Farmdale	7.5	2.0
565	do.	Morton	3.0	1.0
567	NW NW NE, 16-30N-1W, Marshall	Roxana-I	1.0	0.5
568	do.	Roxana-II	1.2	0.6
569	do.	Roxana-III	2.8	2.2
570	do.			
571	do.	do. do.	2.8 2.8	1.6 0.8
572	do.	Roxana-IV	0.8	0.4
573	do.			
574	SE NE NW, 36-20N-7W, Mason	Morton Roxana	5.0 9.5	4.0 7.0
575	do.	do.	9.5	5.0
576	do.	do.	9.5	2.0
577	do.	Peoria	40.0	38.0
578	SE cor., 34-18N-6W, Menard	Roxana	10.0	9.0
579	do.	do.	10.0	4.0
580	do.	Peoria	15.0	12.0
587	SW NE NE, 18-1N-1E, Schuyler	Roxana-I	6.0	5.0
588	do.	do.	6.0	3.0
589	do.	Roxana-II	6.0	4.0
591	do.	do.	6.0	1.0
592	do.			
593		Roxana-III	8.0	6.0
	do.	do.	8.0	1.0
594	do.	Roxana-IV	7.5	3.5
595 596	do. do.	do. Peoria	7.5	0.2
			26.0	24.0
597	do.	do.	26.0	17.0
598	do.	do.	26.0	11.0
599	do.	do.	26.0	8.0
600	do.	do.	26.0	3.0
606	SE NE NW, 16-1N-1W, Schuyler	Roxana	10.0	8.0
607	do.	do.	10.0	4.5
608	do.	do.	10.0	3.0
609	do.	do.	10.0	1.0
610	do.	Peoria	10.0	8.0
617	Center NE, 22-1N-1W, Schuyler	Roxana	5.0	1.0
618	do.	Peoria	12.0	11.0
623	SW SW NW, 23-1N-1W, Schuyler	Roxana	15.5	14.5
624	do.	do.	15.5	10.5
625	do.	do.	15.5	7.0
626	do.	do.	15.5	1.0
626-A				
626-B	do.	do.	15.5	0.5
	do.	Peoria	20.0	19.5
626-C	do.	do.	20.0	19.0
626 - D 627	do.	do.	20.0	18.0
	do.	do.	20.0	18.0
628	do.	do.	20.0	8.0
629	do.	do.	20.0	2.0
633	NW NE NE, 21-13S-1W, Calhoun	Valderan ow.	8.0	6.0
634	do.	do.	8.0	3.0
635	SE SE NE, 28-13S-1W, Calhoun	Altonian ow.	18.0	13.5

TABLE 1 — Continued

Sample no.	Location 1/4 Sec., T., R., County	Stratigraphic unit	Thickness of unit	Feet below top of unit
	V			
637	SE SE NE, 28-13S-1W, Calhoun	Altonian ow.	18.0	6.5
639	do.	do.	18.0	2.5
640	do.	Woodf, ow.	31.5	31.0
642	do.	do.	31.5	27.0
644	do.	do.	31.5	20.5
645	do.	do.	31.5	17.5
647	do.	do.	49.5	8.0
652	SW NE SW, 23-1N-1W, Schuyler	Roxana	25.0	24.5
653	do.	do.	25.0	22.0
654	do.	do.	25.0	17.5
655	do.	do.	25.0	12.0
656	do.	do.	25.0	7.0
657	do.	do.	25.0	1.0
658	do.	Peoria	13.0	12.0
659	do.	do.	13.0	10.0
660	do.	do.	13.0	8.0
661	do.	do.	13.0	5.0
662	do.	do.	13.0	3.0
663	SW SW NW, 23-1N-1W, Schuyler	Peoria	20.0	20.0
664	do.	do.	20.0	16.0
665	do.	do.	20.0	14.0
666	do.	do.	20.0	12.0
667	do.	do.	20.0	10.0
668	do.	do.	20.0	9.0
669	do.	do.	20.0	8.5
670	do.	do.		8.0
671	do.	do.	20.0	
672		-	20.0	7.5
692	do.	do.	20.0	6.5
693	Center, 31-26N-3W, Tazewell do.	Roxana Farmdale	0.8 3.5	0.4
		raimuale	3.5	3.2
694	do.	do.	3.5	2.0
695	do.	do.	3.5	0.5
719	NW SW SE, 9-28N-6E, Stephenson	Roxana	1.0	0.5
720	do.	P e oria	3.0	2.0
723	SE SW NW, 8-23N-9E, Ogle	Morton	1.3	0.7
743	NW SW NW, 1-19N-12W, Vermilion	Shelbyville ow.	5.0	2.0
761	NW NW, 10-24N-5E, Carroll	Peoria	6.0	5.0
776	NE cor. NW, 3-14N-4W, Sangamon	Peoria	5.0	4.5
790	NW SW SW, 8-7N-3E, Fulton	Peoria	3.5	3.0
801	SE NW NE, 29-3N-8W, Madison	Peoria	71.5	60.0
802	do.	do.	71.5	38.0
803	do.	do.	71.5	37.5
804	do.	do.	71.5	35.5
805	do.	do.	71.5	35.0
823	SW NE, 21-33N-5E, LaSalle	Woodf. ow.	8.0	7.0
824	do.	do.	8.0	3.0
825	do.	do.	8.0	1.0
827	do.	do.	20.0	16.0
021				
827 829	SW SE SE, 15-33N-3E, LaSalle	Woodf. ow.	20.0	1.0

TABLE 1 - Continued

Sample no.	Location 1/4 Sec., T., R., County	Stratigraphic unit	Thickness of unit	Feet below top of unit
831	SW SE SE, 15-33N-3E, LaSalle	Woodf. ow.	1.0	0.3
832	SE SW NE, 35-16N-11E, Bureau	do. (sil	t bed) 1.8	1.4
833	do.	do.	do. 1.8	1.0
834	do.	do.	do. 1.8	0.8
835	do.	do.	do. 1.8	0.2
836	do.	do.	35.0	13.0
845	SW NE SE, 17-11N-9E, Peoria	do.	70.0	20.0
846	do.	do.	70.0	8.0
847	NE SE NE, 10-6N-5E, Fulton	Roxana	7.0	0.3
848	do.	Farmdale	0.5	0.2
849	do.	Peoria	18.0	17.8
850	do.	do.	18.0	17.3
851	do.	do.	18.0	16.8
852	do.	do.	18.0	16.3
853	do.	do.	18.0	15.8
854	do.	do.	18.0	15.0
855	do.	do.	18.0	14.0
857	NE SE SW, 2-4N-3E, Fulton	Roxana	8.5	7.5
858	do.	do.	8.5	6.5
859	do.	do.	8.5	5.5
860	do.	do.	8.5	4.0
861	do.	do.	8.5	2.5
862	do.	do.	8.5	1.5
863 864	do. do.	do. Farmdale	8.5 1.5	0.5 1.0
865 866	do. do.	do. Peoria	1.5 17.0	0.5 17.0
867	do.	do.	17.0	16.5
868	do.	do.	17.0	16.0
869	do.	do.	17.0	15.5
870	do.	do.	17.0	15.0
871	do.	do.	17.0	14.5
872	do.	do.	17.0	14.0
873	do.	do.	17.0	12.5
874	do.	do.	17.0	11.5
875	do.	do.	17.0	10.5
876	do.	do.	17.0	9.0
877	do.	do.	17.0	8.5
878	do.	do.	17.0	8.0
879	do.	do.	17.0	7.0
880	do.	do.	17.0	5.5
881	do.	do.	17.0	4.0
882	do.	do.	17.0	3.0
883	do.	do.	17.0	2.0
884	do.	do.	17.0	1.0
885	NW NW NE, 11-4N-3E, Fulton	Woodf. ow.	22.0	18.0
886	do.	do.	22.0	13.0
887	do.	do.	22.0	9.0
888	do.	do.	22.0	5.0
889	do.	do.	22.0	2.0

TABLE 1 - Continued

Sample no.	Location 1/4 Sec., T., R., County	Stratigraphic unit	Thickness of unit	Feet below top of unit
890	NW NW NE, 11-4N-3E, Fulton	Peoria	13.0	10.0
891	do.	do.	13.0	8.0
892	do.	do.	13.0	5.0
893	NE SE, 27-14N-13W, Scott	Roxana-I	7.0	6.5
894	do.	do.	7.0	4.0
895	do.	do.	7.0	1.5
896	do.	Roxana-II	6.0	5.0
897	do.	do.	6.0	3.0
898	do.	Roxana-IV	6.0	2.7
899	do.	do.	6.0	1.7
900	do.	do.	6.0	1.2
901	do.	do.	6.0	0.7
902	do.	do.	6.0	0.2
903	do.	Peoria	9.0	8.8
904	do.	do.	9.0	8.3
905	do.	do.	9.0	7.8
	do.	do.	9.0	7.3
906 907		do.	9.0	6.8
	do.		9.0	6.3
908 909	do. do.	do. do.	9.0	5.3
910	do.	do.	9.0	4.3
911	do.	do.	9.0	2.3
913	SW SW NE, 27-12N-13W, Greene	Roxana-I	6.0	5.0
914	do.	do.	6.0	3.5
915	do.	do.	6.0	0.5
916	do₊	Roxana-II	3.5	3.0
917	do.	do.	3.5	1.7
918	do.	do.	3.5	0.5
919	do.	Roxana-III	7.0	6.5
920	do.	do.	7.0	4.5
921	do.	do.	7.0	2.5
922	do.	do.	7.0	0.5
923	do.	Roxana-IV	3.5	3.5
924	do.	do.	3.5	2.5
925	do.	do.	3.5	1.5
926	do.	do.	3.5	0.5
927	do₊	Peoria	12.0	11.5
928	do.	do.	12.0	10.0
929	do.	do.	12.0	8.5
930	do.	do.	12.0	7.0
931	do.	do.	12.0	5.5
932	do.	do.	12.0	4.0
933	do.	do.	12.0	3.0
934	SW NE NE, 28-10N-13W, Greene	Roxana	3.0	2.5
935	do.	do.	3.0	1.0
936	do.	do.	3.0	0.5
930 937	do.	Peoria	30.0	29.5
938	do.	do.	30.0	28.5
/00	uv•			
939	do.	do.	30.0	26.5

TABLE 1 - Continued

Sample no.	Location 1/4 Sec., T., R., County	Stratigraphic unit	Thickness of unit	Feet below top of unit
941	SW NE NE. 28-10N-13W. Greene	Peoria	30.0	20.0
942	do.	do.	30.0	16.0
943	do.	do.	30.0	12.0
944	do.	do.	30.0	10.0
955	SE cor., 26-2N-9W, Adams	Roxana	4.0	2.5
956	do.	Peoria	22.0	20.0
957	do.	100114	22.0	17.0
958	do.	do.	22.0	7.0
959	NW NE SW, 4-3N-8W, Madison	Peoria	20.0	10.0
960	do.	do.	20.0	6.0
961	do.	do.	20.0	3.0
975	NW NE SE, 16-12N-1W, Christian	Roxana	2.0	1.0
976	do.	Peoria	4.0	3.0
982	SW SW SW, 16-8N-1E, Fayette	Roxana	2.5	2.0
983	do.	do.	2.5	1.0
984	SE NW NE, 29-3N-8W, Madison	Peoria	71.5	71.0
985	do.	do.	71.5	68.5
9 86	do.	do.	71.5	66.0
987	do.	do.	71.5	62.0
988	do.	do.	71.5	56.5
989	do.	do.	71.5	52.0
990	do.	do.	71.5	45.0
991	do.	do.	71.5	40.0
992	do.	do.	71.5	35.5
993	do.	do.	71.5	32.0
994	do.	do.	71.5	27.0
995	do.	do.	71.5	22.0
996	do.	do.	71.5	17.0
997	do.	do.	71.5	12.0
998	do.	do.	71.5	6.0
999	NE NW NW, 25-2N-9W, St. Clair	Roxana-I	10.0	7.0
999-A	do.	do.	10.0	10.0
1000	do.	do.	10.0	5.0
1001	do.	do.	10.0	3.0
1002	do.	do.	10.0	1.5
1003	do.	do	10.0	0.5
1004	do.	Roxana-II	12.0	11.0
1005	do.	do.	12.0	8.0
1006	do.	do.	12.0	6.0
1007	do.	do.	12.0	4.0
1008	do.	do.	12.0	2.0
1009	do.	do.	12.0	0.5
1010	do.	Roxana-III	7.0	6.5
1011	do.	do.	7.0	5.0
1012	do.	do.	7.0	3.5
1013	do. do.	do. do.	7.0 7.0	2.0 0.5
1014		do. Roxana-IV	14.0	13.0
1015	do. do.	koxana-iv do.	14.0	9.0
1016		do.		5.0
1017	do.	uo.	14.0	5.0

TABLE 1 - Continued

Sample no.	Location St 1/4 Sec., T., R., County	ratigraphic unit	Thickness of unit	Feet below top of unit
1018 1019	NE NW NW, 25-2N-9W, St. Clair Ro	oxana-IV do.	14.0 14.0	3.0 1.0
1020	do. P∈	eoria	25.5	25.0
1021	₫o.	do.	25.5	23.5
1022	do.	do.	25.5	22.5
1023	do.	do.	25.5	18.0
1024	do.	do.	25.5	14.0
1025	do.	do.	25.5	10.0
1026	do.	do.	25.5	6.0
1028	SW SE SW, 7-6S-5W, Randolph Ro	oxana	5.5	4.8
1029	do.	do.	5.5	2.0
1030		eoria _	4.0	2.0
1048		xana-I	2.0	1.5
1049	do.	do.	2.0	0.5
1050		oxana	2.5	2.4
1051	do.	do.	2.5	2.0
1052	do.	do.	2.5	0.5
1053	•	eoria	6.5	6.0
1054	do.	do.	6.5	4.5
1055	do.	do.	6.5	3.5
1056	do.	do.	6.5	2.0
1059	Interstate 57, 8-11S-2E, Johnson Ro		2.5	1.2
1060		eoria	5.0	3.5
1071		oxana	3.5	2.0
1072	do. Pe	eoria	8.0	6.0
1078		oxana	6.0	5.0
1079	do.	do.	6.0	3.0
1080	do.	do.	6.0	1.0
1081		eoria (trans.)	5.0	4.0
1081-A	do.	do.	5.0	2.5
1082	do.	do.	5.0	0.5
1083		eoria	20.0	18.0
1084	do.	do.	20.0	16.0
1085	do.	do.	20.0	13.0
1086	do.	do.	20.0	11.0
1087	do.	do.	20.0	9.0
1088	do.	do.	20.0	7.0
1089	do.	do.	20.0	5.0
1090 1091	do. do.	do. do.	20.0 20.0	3.0 1.0
1096		oxana	2.0	1.0
1097 1098		eoria (trans.)	2.0 10.0	1.0 9.0
1098	do.	eoria do.	10.0	5.0
1101	NE cor., Henderson, Henderson, Ky.Pe		35.0	34.0
1102	•	eoria	35.0	32.5
1102	do.	do.	35.0 35.0	31.0
1103	do.	do.	35.0	28.0
1105	do.	do.	35.0	25.0
1106	do.	do.	35.0	22.0
****	uo.	u.,	JU . U	22.0

TABLE 1 - Continued

Sample no.	Location 1/4 Sec., T., R., County	Stratigraphic unit	Thickness of unit	Feet below top of unit
1107	NE cor., Henderson, Henderson, Ky.	Peoria	35.0	17.0
1108	do.	do.	35.0	12.0
1110	1.1 mi. W. Green R. br., Henderson, Ky.	Peoria	10.0	6.0
1113	2 mi. W. Owensboro, Davies, Ky.	Peoria	15.0	12.5
1114	E SE of Lancaster Cem., Davies, Ky.	Peoria	14.0	14.0
1115	do.	do.	14.0	11.5
1116	do.	do.	14.0	9.5
1117	do.	do.	14.0	7.5
1118	do.	do.	14.0	6.0
1119	do.	do.	14.0	4.0
1122	NW SW NE, 10-7S-8W, Warrick, Ind.	Roxana	5.0	3.0
1123	do.	do.	5.0	1.0
1124	do.	Peoria	10.0	7.0
1125	do.	do.	10.0	3.0
1129	N_2^1 , 25-6S-9W, Warrick, Ind.	Roxana	3.0	2.0
1130	do.	Peoria (trans.)	1.0	0.5
1131	do.	Peoria	8.0	7.0
1132	do.	do.	8.0	4.0
1135	SE NW SW, 1-5S-14W, Posey, Ind.	Roxana-I	1.5	0.8
1136	do.	Roxana-II-IV	4.5	4.0
1137	do.	do.	4.5	2.5
1138	do.	do.	4.5	0.5
1139	do.	Peoria (trans.)	1.5	0.5
1140	do.	Peoria	34.0	29.0
1141	do.	do.	34.0	25.5
1142	do.	do.	34.0	22.0
1143	do.	do.	34.0	19.0
1144	do.	do.	34.0	16.0
1145	do.	do.	34.0	13.0
1146	do.	do.	34.0	10.0
1147	do.	do.	34.0	6.0
1148	do.	do.	34.0	2.0
1149	NE NE SE, 29-2S-14W, Edwards	Peoria	6.0	3.0
1150	do.	do.	6.0	5.0
1154	NE NW NE, 8-1S-12W, Wabash	Roxana	1.5	0.5
1155	do.	Peoria	6.0	5.0
1156	do.	do.	6.0	2.0
1159	SW NE SW, 6-5N-11W, Crawford	Roxana	1.0	0.5
1160	do.	Peoria	4.0	2.5
1168	NE NE NW, 34-7N-10E, Jasper	Peoria	5.0	2.5
1171	NE NW SW, 10-9N-9E, Cumberland	Peoria	2.0	1.0
1173	SW SW SW, 23-9N-8E, Cumberland	Peoria	3.0	2.0
1176	NW NW NW, 30-8N-4E, Effingham	Peoria	5.0	2.0
1178	SE SE NE, 24-8N-3E, Fayette	Peoria	6.0	4.0
1182	SE SE NE, 21-11N-3E, Shelby	Roxana	2.0	1.0
1183	do.	Peoria	3.0	1.5
1184	SW SE NE, 26-5S-6W, Pike	Roxana	3.0	1.5
1185	do.	Peoria	5.0	2.5
1191	NW NE SE, 9-5S-6W, Pike	Roxana	3.5	2.0
1192	do.	Peoria	7.5	6.8

TABLE 1 - Continued

Sample	Location	Stratigraphic	Thickness	Feet below
no.	1/4 Sec., T., R., County	unit	of unit	top of unit
1193	NW NE SE, 9-5S-6W, Pike	Peoria	7.5	5.0
1194	NW SW SW, 32-35-7W, Adams	Peoria	8.0	8.0
1197	SW cor., 32-3S-7W, Adams	Roxana	4.0	2.0
1198	do.	Peoria	15.0	13.0
1199	do.	do.	15.0	10.0
1214	NW NE NW, 23-15-9W, Adams	Roxana	4.0	0.5
1215	do.	Peoria	37.0	36.0
1216	do.	do.	37.0	34.5
1217	do.	do.	37.0	32.5
1218	do.	do.	37.0	30.0
1219	do.	do.	37.0	26.0
1220	do.	do.	37.0	22.0
1221	do.	.do.	37.0	18.0
1222	do.	do.	37.0	15.0
1223	do.	do.	37.0	10.0
1227	SE cor., 26-2N-9W, Adams	Roxana	4.0	3.5
1228	do.	do.	4.0	1.5
1229	do.	do.	4.0	0.5
1230	do.	Peoria	22.0	20.0
1231	do.	do.	22.0	17.0
1232	do.	do.	22.0	13.0
1233	do.	do.	22.0	9.0
1234	do.	do.	22.0	6.0
1235	Center E_2^1 , 31-5N-8W, Hancock	Roxana	3.0	1.5
1236	do.	Peoria	6.0	4.0
1239	SW NE SW, 36-8N-7W, Henderson	Roxana	4.0	3.5
1240	do.	do.	4.0	2.0
1241	do.	do.	4.0	1.0
1242	do.	Peoria	31.5	30.5
1243	do.	Peoria	31.5	29.5
1244	do.	do.	31.5	26.0
1245	do.	do.	31.5	23.0
1246	do.	do.	31.5	21.0
1247	do.	do.	31.5	17.0
1248	do.	do.	31.5	11.0

^{*} ow. = outwash

TABLE 2—HEAVY AND LIGHT MINERAL ANALYSES

		Percent soluble	15.4	23.6	21.8	10.6	1	1	•	21.6	9.2	11.0	8.6	•	•	ı	•	1	ı	1	1		17.5	12.5	18.5	4.5	4.0	4.0 0.0
		Офрега	,	ı	,	,			ı	,	ŧ	ı	ı	-	ı	1	ı	Ø	1	N	က	ო	ď	۱ (0	-	ı m	4 ω
	(%)	.ebiei ea-6N		1	1	ı	:		ı	ı	ı	ı		4	ı	7	1	7	1	9	7	Ŋ	٧	' '	17	9	2	8 13
	Light	K-feldspar		1	•	ı	ı	1	•	,	ŧ	ı	ı	12	1	33	ı	16	•	16	14	7	17	. 1	Ŋ	16	72	15 14
	ı	Straup	١.	ı	1	•	•	ı	1	1	i	ı	•	83	1	2	ı	75	ŧ	92	75	82	72	! !	99	77	99	73
		Офрега		6	ω	7	1	1	1	-	23	7	ო	-	1	ı	•	1	ı	ŧ	ı	7	C	1 4	ß	N	4	Ø 10
		•BuA-•qoid	ŀ	4	_	ო	ı	1	•	1	1	4	0	N	1	_	0	ო	4	4	7	•	7	ا	_	ı	ı	: -
		Hypersthene	,	0	-	-	1	ı	ı	1	_	-	ო	•	1	•	ı	•	,	ı	1	•	ď	4	ı	ī	ı	1 (1)
		Enstatite	١	•	1	t	t	ო	0	1	ı	i	•	ഹ	က	4	7	-	0	4	4	ı	-	. –	1	•	1	1 0
	(%)	Hornblende	23	27	36	21	22	35	56	32	7	36	17	23	35	38	24	8	22	33	39	36	42	် လိ	99	75	5,4	33
		Actinolite	5	ო	2	2	-	1	4	9	6	7	4	1	-	-	7	~	വ	7	ı	വ	0	u ا		ı	•	13
1s)	Minerals	Sillimanitie	١,	ı	ı	ı	ı	ı	1	ı	ı	ı	•	1	1	•	ı	ı	ı	ı	ı	ı	ı	ı	ı	ı	ı	1 1
basis)	y Mi	Andalusite	-	-	ო	-	ı	ı	1	ı	α	1	ı	ı	ı	1	ı	7	ı		ı	٦	ı	•	ı	ı	•	1 1
free	Heavy	Staurolite	١,	ı	ı	1	ı	ı	_	ı	1	•		-	ı	ı	ı	ı	0	ı	-	1	_	٠	-		1	1 1
ate-		Kyanite	١,	ı	1	ı	1	ŧ	ı	ı	ŧ	1	:	ı	1	,	ı	ı	ı	1	ı	ŧ	1	1	ı	ı	ı	1 1
(carbonate-free	Transparent	Piranite	١,	,	-	ı	ı	ı	-	į	•	7	7		ŧ	,		f	ı	•	ı	ı	,		1	ı	•	٠ -
(cai	Fran	AlituA	က	~	,	1	7	ı	_	-	1	٦	-	í	•		~	ı	ı	ı	ı	ı	ı	-		ı	1	1 1
	-	Epidote	22	17	17	24	36	33	33	ട്ട	13	13	53	41	6	36	44	34	36	23	34	41	6		13	13	52	28 26
		Garnet	0		7					16	10	11	15	13					6				9	: =	6	ഹ	10	18
		Zircon	23	20	7	22	9	20	22	13	17	4	17	2	ω	ω	11	ω	14	20	11	٦	_	ı Ω	က	വ	10	16
		enilamzuoT	4	N	9	-	ო	ı	1	9	4	11	7	က	-	-	~	ო	ო	7	ო	-	m	2	N			വ വ
	dne	Офрега		0	24		ო	16	13	6	7	53	40	ω	12	12	11	10	13	12	15	ω	4	6	Ŋ	7	12	13
	Opaque	ВЈзск	ြိ	ഗ	CA	ഗ	മ	15	2	4	(1)	CA	4	19	8	19	54	19	22	27	35	21	32	8	വ	œ	38	30
		Heavy minerals in fraction (%)	3	ı	1	i	ı	1	1	1	ı	1	1	1	1	1	•	1	1	•	•	ı	2.1	2.4	1.2	ı	1.3	8°°°
		.062- .250mm fraction (%)	7.6																									30.1
		Sample no•	8	10	11	12	17	18	19	20	21	22	23	೫	31	32	33	8	32	36	36	40	76	77	83	84	82	86 87

		Percent soluble	2.5 5.5 5.0	4.5 18.0 19.5 11.5	3.5 16.5 40.0 21.0 39.0	30.0 0.0 0.0 0.0	86.0 30.5 30.5	31.0 40.5 41.5
		Others	ოოა	10 10 3	п.400	ωιιιι	വവവവ	4 m r
	(%	wa-Ca felds.	8 9 16	8 11 8 7	7 7 13 10	01111	7 10 8 8	9 6 91
	Light	K-feldspar	20 13 16	23 10 18 18	22 - 18 13 19	15	14 17 20 15 18	17 9 14
	L	StieuQ	70 75 63	60 72 77 72	70 - 71 68 65	2	75 73 70 70	5 5 5 8
		Офрега	ι ю ι	41011	11701	10101	ω ο ι ι ω	011
		•QuA-•qoid	ю О I	1 -4 1 1 1	4 4 1 2 4	11400	1 1 1 1 1	1 1
		Hypersthene	- 1 1	1 1 1 1 1	1 1 1 1 1	1 1 1 1 1	1 1 1 1 1	1 1 1
		Enstatite	-2-	агю гъ	1 4 1 4 1	ч п п 4 ю	1 1 1 1 1	ιιω
	(%)	Hornblende	36 32 52	53 43 35 35	33 33 50 55 55	84 4 4 6 8 8 4 4 8 8 8 8 8 8 8 8 8 8 8 8	29 55 63 67	65 61 38
ъ	als	Actinolite	ω 4 Ω	17158	n n 2 a u		1 1 1 1 1	1 1 -
TABLE 2—Continued	Minerals	Sillimanite	1 1 1	1 1 1 1 1	3 1 1 1 1	1 1 1 1 1	1 1 1 1 1	1 1 1
Cont		ArisulsbnA	1 1 1	1 1 1 1 1	t 1 1 1 t	1 1 1 1 1	1 1 1 1 1	1 1 1
2—(Неаvу	Staurolite	110			1111	1 1 1 1 1	1 1 1
BLE	Transparent	Kyanite	1 1 1	1 1 1 1 1		1 1 1 1 1	1 1 1 1 1	1 1 1
TA	spa	Titanite	 1 1	11011	4 4 4 4 1	11 1 2 2		1 1 1
	Tran	AlituR	1 - 1	1 . 1 1 1			1 1 1 1 1	1 1 1
		Epidote	38 32 22 22	88 98 88 88 88 88	23 23 24 24 24	36 31 34 38 47	32 17 17 11	17
		fearnet	7 14 14	4 91 8 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4 4	15 7 11 6	6 15 10 8 17	15 19 13 12	23 8 9
		Zircon	900	15 15 15	11 7 12 10	333	13 8 9 6	6 31 31
		Fourmaline	040	детее	- 1 - 1 1	10011		4 1 2
	aque	Others	7 7	7 10 10 13	8 8 7 9 13	10 20 17 14	13	11 12 11
	Opa	ВЈЗСК	23 10 13	8 4 5 4 5 8 3 3 4 5 4 5 8	33 34 34 35	23 20 15 15	22222	25 7 4
		Heavy minerals in fraction (%)	4.4 0.7 0.7	0 0 0 4 1 4 5 1 4	1111	1 1 1 1 1	0 1 1 1 1	1 . 1
•		.062- .250mm fraction (%)	1.8 3.6 4.0	44 10 40 5.	446.00	30.2 13.2 0.2 0.2	13. 4.00 4.4. 4.00	5.2 - 11.5
ļ		Sample no.	88 86 66	100 101 103 103	105 106 107 109	110 115 116 117	133 134 135 136 137	140 144 147

																													33
		Percent soluble	ı	t	ı	ı	•	ı	ı	ı	•	•	•	•	1	1	1	11.5	28.0	4.5	7.5	4.0	5.5	2.5	4.0	15.5	7.5	19.0	4.5
		Оґрега	ı	ı	ı	ı	ı	•	1	i	•	ı	ı	ı	ı		ı	ı	ı	ı	ı	1	4	4	ı	ı	ı	•	က
	(%)	.sbieł &S-sN	,	ı	ŧ	ı	ı	ı	1	•	ı	ı	ı	ı	ı	ı	ı	ı	1	,	ı	ı	12	11	,	ı	;	ı	10
	Light	K-feldspar	1	,	ı	ı	1	ı	1	1	ı	i	ı	ı	•	ı	î	ı	•	ı	ı	ı	16	10	ı	ı	ı	1	Π
	T	Cuartz	,	,	ı	ı	ŧ		,	1	ı	•	ı	1	•	ı	1	ı	ı		1		89	75	ı	!	ı	ı	92
		Ofpers	ಬ	ო	9	~	-	-	-	4	f	4	ı	7	ı	-	ო	ო	ı	7	8	-	ı	ო	-	ო	ო	7	7
		•guA-•qoid		ı	_	-	-	ч	7	10	12	11	ო	N	_	ı	ı	ო	1	0	-	ŧ	ı	-	ı	1	ı	ı	•
		Hypersthene	,	,	ī	ო	က	4	12	6	16	14	ო	•	ſ	Ŋ	ı	ı	1	Ø	9	_	1	0	ı	1	ı	ı	0
		Enstatite	1	_	-	8	_	-						വ	-		ო	-	-	1	ı	-	ŧ	0	-	•	•	-	1
	(%)	Hornblende	77	72	65	63	5	53	36	56	88	27	46	44	53	36	24	36	41	3	57	8	4	9/	41	48	48	74	99
	Minerals	Actinolite	,		ı	0	ı	ı	•	-	α	ı	ო	~	~	ı	~		ო	ı	ı	ı	ı	ı	ı	1	ı	ı	ო
	Mine	Sillimanite		1	ı	ı	ı	ı	ı	ı	•	ı	1	ı	ı	ı	ı	,	•	ı	ı	ı	ı	1	ı	ı	ı	ı	,
	Неаvу	Andalusite	١	١	1	ŧ	;	١	١	,	ı	ı	•	1	•	1	ı	ı	•	ı	1	ι	1	ı	٠	1	1	•	t
		Staurolite	ŧ	١	1	ı	1	1	ı	ı	1	,	1	-	1	ı	ı	ı	ı	1	•	•	:	7	ı	ı	ŧ	ı	•
	rent	Kyanite	•	ı	1	ı	ı	ı	,	ı	ı	,	ı	ı	•	•	1	t	ı	ı	ı	1	ı	1	ı	ı	ı	•	1
	Transparent	Titanite	1	-	-	١	~	ı	1	-	ı	1	1	•	~	ı	•	ო	i	1	•	1	1	ł	1	1	1	•	1
		Rutile	7	1	ı	ı	ī	-	1	1	•	1	1	~	ı	1	1	1	t	1	•	•	1	ı	ı		ı	ı	•
		Epidote	9	13	10	17	10	ω	18	15	22	12	32	89	39	43	52	32	43	ß	11	54	33	4	16	22	9	6	ω
		Garnet	æ	6	10	ω	19	22	19	31	17	8	11	Ŋ	7	11	14	18	Ŋ	20	20	6	10	7	ω	15	9	11	16
		Nooris	7	-	٦	0	10	9	വ		က	4	٦	-	9	0	0	~	ო	34	0	88	12	7	33	6	37	Ø	N
		Tourmaline	ł	1	٠	ı	ı	1	4	ო	ı	ı	~	ı	-	0	-	7	4	9		-	-	ო	ı	ო	1	~	~
	dne	Офрега	9	8	15	9	11	16	13	10	ω	19	11	σ	13	33	14	വ	ი	16	6	13	12	19	20	4	17	17	11
	Opa	ВЈ9СК	36	23	54	വ	18	2	17	41	39	ထ္တ	14	24	77	24	7	15	23	19	11	8	32	6	ဓင္တ	42	20	6	Ŋ
		Heavy minerals in fraction (%)	8.3	1.9	5.9	1.0	2.8	2.0	3,3	6.2	4.0	7.2	3.0	3.6	6.4	4.3	3.9	1.5	0.5	0.7	1.5	2.6	0.3	1.4	1,5	1.1	6.0	1.8	0.7
		.062- .250mm fraction (%)	3.6	2.4	2.4	27.6	18.9	36.5	33•3	9.5	8.9	12,5	20.3	2.4	8.7	9.5	7.2	17.2	3.2	32.7	46.3	5.4	5.2	25.5	28.9	17.7	24.2	26.3	28.6
ļ		Sample no.	152	153	154	155	156	157	158	159	160	161	162	163	164	165	166	175	177	178	179	271	272	280	281	282	283	285	286

TABLE 2—Continued

		Percent soluble	19.0 25.5 4.5 2.0	27.0 36.5	33.0	13.0 46.5 35.5 7.5	7.0 5.5 17.0 52.5 2.0	10.0 8.0 9.5 9.0	16.0 16.5
	_	Others	13 6 8 8 8	1 000-4	0 4	212 21 3 1 3 1 3 1	7767	99999	iO I
	(%)	Na-Ca felds.	7 8 9 9 8	10 10	ν ω	65.13	ε α 4 7 6 9	4 ~ 9 7 ~	14 6
	Light	K-feldspar	112	111	12	111 17 100 100 100 100 100 100 100 100 1	15 15 18 12	13 11 15 15	15 15
		ZireuQ	67 72 72 74	8446	3,5	84 51 77 83	80 75 62 67 72	81 80 79 68 76	99
		Офрека	110011	1144	t	וומוה	01500	1 - 1 0	I
		•Bn¥-•doid		1 1 0	n 01	01001	- 0 - 0 -	1114	ч т
		Hypersthene		1	9	r 1 w m 1 -1	1 01 10 1	1 1 1 1	1 1
	_	Enstatite			o	w + 01 - 1	ოიოო၊	1 1 1 1	7 7
	88	Hornblende	63 60 60 43 37 51	27.22	62	62 174 34 34	61 66 72 50 33	38 15 31 56 48	51
	rals	Actinolite	HIIMM	111	-		Llaga	1 1 1 1	1 1
	Mineral	Sillimanilis	1 1 1 1 1						
	Heavy	Andalusite	1 1 1 1 1		i				
5		Staurolite	1111		٠,			1111	
	rent	Kyanite Kyanite			1				1 1
	Transparent	Titanite							
	Tra	AlituA			,	11111		1011	11
		Epidote	844 84 84 84 84 84 84 84 84 84 84 84 84	36 13 7	17	61766	21 15 15 34	18 75 8 8 8 8 1 8 8 1 8 8 1 8 8 1 8 1 8 1 8	11
		Garnet	6 9 7 7		- 0	9 - 11 - 9 13 13	10 B C C C C C C C C C C C C C C C C C C	78217	13
			444000			32-2-5	01110	25 31 10 10 26	6 2
		Zircon		1 1 6 -		01414	010	PHH 10	
	e)	Tourmaline	17 15 20 12	33		16 - 27 13 18	11 24 24	36 22 21 27 17	၈ မှ
	aque	Офрек			- 10				3 1
	o d	BJ9ck	019472	12 61	•	51.634	100	4 % · ·	11
		Heavy minerals in fraction (%)	3.2 trace 0.4	0 0 0 0 0 0 4 0	1.4	0.000	1 2 0 0 0 1 1 1 0 0 0 0 1 1 1 1 1 1 1 1	0.1 0.4 0.0 0.0	00.0
		.062- .250mm fraction (%)	6.4.8 8.0.1 0.01	1.2 2.0 6.6	21.9	34.2 1.1 3.6 7.8 7.8	14.5 40.9 26.9 8.9	4.8 23.0 20.5 11.8 16.1	11.6 8.5
		Sample no.	310 311 332 333 341	342 351 353	366	387 389 406 414 425	4 4 4 4 4 4 4 4 4 4 3 3 4 4 3 5 4 4 3 5 4 4 3 5 4 4 3 5 4 5 5 4 5 6 6 6 6 6 6 6 6 6 6 6 6 6 6	460 461 462 464	465-A 466

Parket depotation or an article for the	Percent soluble	8.0 11.0 30.0	22.0	ο α Ο ι	29.5	25.5	8	25.0 21.5	9.0	21.0	ဂ ဂ	21.0	11.0	35.0	8.0 9.5 7.5
_	Офрета	600	ľΩ	V (V (n	0.4		1 0	1 1	ı m	_	4 ı	; ;	ı	1110
8	.abiel so-sN	9 6 19	σ (, w	0 ~	8 2	I	ı œ	1 1	12	~	= -		r	11100
Light	K-feldspar	13 14 15	12	177	13	12	1	1 02	1 1	133	, T	9 1		1	1 1 23
ä	Quartz	75 78 63	74	281	υ δ	78	. ;	- 02	f I	- 62	ر د	60 1		1	1119
	Офрета	418	4	1		⊣ "		ეფ	ω 4	0 00		2 -	1 4	2	0001
	•bn A-• qoid		F 1	1	, ,	4 1	3	1 m	1 0	1 -	-	I	i	4	4401
	Нуретѕthепе	1 1 1	1 1			I	ı	i m	1 ~	- 0	-	N 1	1 1	•	
	Enstatite	1 1 1				4 1	i	7 7	1 4		N				1101
8	Hornblende	45 42 33	31	881	21	57	26	62 55	49 57	54 54	4	65	3	73	68 70 72 61
als	Actinolite	1 1 1		1 1	: :	ហ៖		ო	ıσ	1 1	ı	⊣ ෆ	ı e	4	1101
Mineral	Sillimanitie	1 1 1	1 8	1 1	: I	1 1	ı	1 1	1 1	1 1	,	1 1		i	1 1 1 1
	ejisu[sbnA			. 1	1 1		,	1 1	1 1	1 1	1	1 1	; i	1	1 1 1 1
Heavy	Staurolite	1 1 1	1 0	n 01	1 1	1 1	1	1 1	0 1		ı	1 1	1 1	1	1110
rent	Kyanite	+ + +	t I	t	1 1	1 1	1		١ -	٦ ١	ı	1 1	1 1	ı	1 1 1 1
Transparent	Titanite	111	1 1	ı t	1 1				1 1	, ,	ı	1	ı ı	ı	
Irai	Rutile	111	1,	- 1	1 1	1 1	-	- •	1) 1	-	1 1	۱ ب	i	
•	Epidote	488	28	98	6	ω ι	14	16 8	31	34	<u> </u>	17	27 2	စ္	111
	Garnet	9 7 13	19	12	31	17	18	12	4 ω	មា	00	7 7	15	7	۲ e c و 4
	Zircon	24 19	18	\$ 23	11	ო 1	ო	က၊	9 6		ı		ი -	-	9599
	PnilsmruoT	001	1 (N (N 11	٠.		1 1	4 ω	H 04	N	⊣ ઌ	C) i	8424
dne	Офрега	17	9 9	17	4 4	12	12	8 %	14 24	16	77.	27	14	2 2	27 28 11
Opa	ВІзск	282	32	261	71	90	11	დო	788	91	Ω	7	3 28	သူထ	6 12 14
	Heavy minerals in fraction	0.6													0.5
	.062- .250mm Sample fraction no. (%)	17.3 12.4 12.4	6.7	12.8	5.2 11.9	8.3	3,4	0.6 21.6	1.1	0 0 4 6	F • 9	1.1 21.5	က လ လ လ	0.1	8 0 1 8 0 3 3 4 4 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5
	Sample	467 468 470	471	473	478 479	480 481	483	484 494	504 505	525 526	532	536 544	545 547	557	568 569 572 574

TABLE 2—Continued

	Percent soluble	32.5	11.5	11.5	0 0	9.0	4.5	15.0	58.5	11.5	38.0		•	•	ı	1	1	•	•	,	1	•	•	•	•
	Others		•	1	1 1		9	4	12	N	7	ı	ŧ	r	,	1	•	ı	1	ı	ŧ	•	ŧ	ı	•
8	Na-Ca felds.	1	,	•		ı	7	7	17	4	6	ı	ŧ	ł	•	ı	•	ı	ı	ı	ı	ŧ	ı	,	ł
Light	K-feldspar	•	1			. 1	20	81	24	12	18					1		•	ı		,	,			ı
:1	Quartz	1								82						ı		•	ı	ï	1			ı	,
\vdash	Офреть	3		0	n -	4		,	1	ŀ	1	0	ო	_	ı	ı	~	ı	ო	10	Q	-	ı	ı	
	•BuA-•qoid	က		α,		t 0	-	7	1	,	,	7	N		ı	1	ı	,	i	ı	ı	-	ı	ı	
	Hypersthene	7	1		ıc	٦ -	വ	-	,	٦	,	N	9	ı	-	٦	1	,	-	7	7	-	1	ı	
	Enstatite	S	-	7		-	1	-	1	ı	ı	1	_	•	0	-	ı	ı	١	ı	-	ı	ı	1	
8	Hornblende	89	9	92 9	200	88	61	28	,	49	4	2	73	62	33	38	36	45	69	20	44	65	21	47	26
rals	Actinolite	ო		ı	1 1		ı		•	ı	ı	ı	ı	ı	1	8	1	•	ı	1	ı	,		ŧ	•
Minerals	Sillimanite	ı		ι	• (1	1	1	ı	ı	ı	•	ı	ı	ı	ı	•	1	ı	1	ı	1	•	1	ı
	Andalusite	,	1	•	1 1	1	ı	1	ı	ı	1	1	ı	•	ı		ı	1	ı		•	•		ı	•
Heavy	Staurolite	1	_	•	1	1	-	ł	ı	ı	ı	1	•	1	١	0	•	1	•	ı	ı	•	•	1	ო
ent	Kyanite	ŀ	Ø	ı		1	ı	•	•	•	ı	•	ı	ı	1	•	,	t	•	1	7	٦	ı	ı	-
Spa	Titanite	,	ī	-		1	1	_	1			F	1	•	ı	•	N	•	,	ı	ı	•	•	,	٠
Transparent	Rutile	ı	7	•	• 1	1	ı	ı	•	ı	1	•	-	1	ı	1	-	٠	ł	t	•	1	1	1	-
	Epidote	12	30	ω (သ င	32	15	೪	•	22	σ	4	4	13	8	50	52	22	17	12	14	20	17	19	19
	Garnet	4	7	വ	m o	р (V	15	1	1	8	10	16	7	18	33	32	ß	17	ഹ	ო	35	ω	15	23	15
	NooriS	-	-	•	١ -	- 1	•	~	1	4	ı	0	ı	N	10	ı	ω	14	ო	ო	-	1	7	0	Ø
	Tourmaline		0	ı N	ກດ	o 01	~	4	1	4	7	0	က	ო		4	7	7	N	٦	~	က	σ	ო	7
Opaque	Others	19	13	13	<u>α</u> ς	3 ~	14	19	1	15	13	22	49	0	15	15	6	17	7	19	10	19	8	15	10
o O Ba		7	ω	ۍ ر <u>ي</u>	٠ -	11	12	10	•	15	24	13	-	15	32	21	ဓ္က	21	ω	4	18	7	4	13	17
	Heavy minerals in fraction (%)	trace				0.0						1.3	8.1	1.4	0.4	0.5	0.2	0.2	2.9	1.8	1,3	1.2	8.5	0.7	0.3
	.062- .250mm fraction (%)	0.2	5.8	3.1	7.6	9.0 9.0	56.4	2.0	3.8	20.2	10.4	25.0	1.7	3.4	17.8	9.3	10.2	5.1	6.0	9.0	6.5	6.0	9.0	16.9	15.0
	Sample no•	577	588	589	200	594 594	633	720	725	729	755	818	827	831	828	860	862	864	867	879	882	688	892	913	915

		Percent soluble		1 1 1 1		1 1 1	1 1	1 1 1 1 1	1 1 1
f		Others	11111	1 1 1 1	1111	1 1 1			1 1 1
	8	sbieł 62-6N	. , , , ,	1 1 1 1	1 1 1 1 1		i i	1111	
	Light	K-feldspar	1111	1 1 1 1	1 1 1 1	1 1 1	1 1	1 1 1 1 1	1 1 1
	3	Quartz	1111	1 1 1 1	1 1 1 1	1 1 1	1 1	1 1 1 1 1	1 1 1
f		Others	- 1-0-	4411	10440	ოოი	19	-011-	1 m 1
		•pu A-• qoid		m 1 1 1		1 1 1	1 1		1 1 1
		Hypersthene	11101	1 1 6 1		1 1 1	ı -	11101	ı ı
	:	Enstatite	11111	0111	10111	1 1 1	10	1 1 1 1 1	1 1 1
	8	Hornblende	\$ 525 525 54 55 55 56 56	50 64 70 70 70 70 70 70 70 70 70 70 70 70 70	66 15 22 57 61	59 55 41	34 68 89	63 30 35 57 69	67 56 64
	als	Actinolite		014 I	~	1 1 1	1 1	1 1 1 1 1	1 1 1
	Mineral	etinamilli2	1111	1 1 1 1	1111	1 1 1	1 1	1 1 1 1 1	1 1 1
		etieulsbnA	11111	1 1 1 1	1111	1 1 1	1 1	1 1 1 1 1	1 1 1
	Heavy	Staurolite		1 -1 1	10110	00 1-10	4 1	w & w	- 24
	ent	Kyanite	10111	1444	LIALI	1 1 4	1 1	11011	444
	par	ətinatiT		-11	+ + + 1 + 1 + 1 + 1 + 1 + 1 + 1 + 1 + 1	1 1 1	1 1	1 1 1 1 1	
	Transparent	Rutile	111	1141	14011	1 -1	I	11011	1 1 1
	_	Epidote	27 16 29 17 16	23 7 6	014822	22.24	21 17	25 33 29 15	19 19 17
		Garnet	11 6 9 6 4	6 16 16	24 4 110	6 01 14	27	10 10 9	യ യ യ
		nosiiZ	11 6 11 9 7	84 4·21	622800	r 4 0	∞	25 25 26 4	100
		enilsmiuoT	4	- மஜக-	4414	⊣	ر ا	-000N-	44 I
-	ant	Оґретѕ	28 4 1 1 8 2 5 5 6 5 6 5 6 5 6 5 6 6 6 6 6 6 6 6 6	23 28 12 12	17 15 15 8	4 4 91	4 8	16 18 18 6	7 12 5
	Opaque	Вјаск	42 5 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	21 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5	49 40 113	17 15 32	37	4 33 29 11	14 23 10
		Heavy minerals in fraction (%)	4 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6 6	0.000.3	00000 04464	0.0 4.0 7.0	0.7	00000	0.1
		.062- .250mm fraction (%)	7.9 10.0 3.0 2.0 2.0	0.4.0	0.5 16.0 3.0 2.9	2.0 0.0 4.0	7.6	4.7 18.7 13.9 6.9 11.0	7.5 5.3 16.5
		Sample no.	917 920 924 928	1101 1122 1125 1137	1141 1214 1215 1217 1218	1220 1223 1228	1229 1231	1233 1239 1241 1242 1243	1245 1246 1248

TABLE 3—AVERAGES OF SELECTED HEAVY AND LIGHT MINERALS FOR STRATIGRAPHIC UNITS AND DEPOSITIONAL PROVINCES

		1	т										1		
				(0			cent anal				e 2)		n	Light ninera	
Depositional province and stratigraphic unit	Number of	Black opaque	Tourmaline	Zircon	Garnet	Epidote	Staurolite	Actinolite	Hornblende	Enstatite	Hypersthene	Diopside- augite	K-feldspar	Na-Ca feldspar	Others
Illinois River Valle Peoria Loess Morton Loess Richland Loess Rock River Valley	18 5 3	20 13 14	1 1	8 6 4	8 8 14	22 14 17	0 0 0	1 1 0	55 64 58	1 1 0	tr. 0 1	tr. 1 0	15 14 16	9 8 6	76 78 78
Peoria Loess Mississippi Valley above Alton	4	15	2	9	8	16	0	0	56	1	1	1	15	6	79
Peoria Loess Mississippi Valley below Alton	20	15	2	5	6	23	1	tr.	59	tr.	1	tr.	16	9	75
Peoria Loess Ohio River Valley	11	22	3	15	11	26	tr.	4	33	1	tr.	1	15	11	74
Peoria Loess Wabash River Valley	3	13	3	5	19	13	0	3	50	1	2	1	13	8	79
Peoria Loess	1	4	1	3	12	10	0	7	66	0	1	0	-	-	-
Illinois Valley (upper) Farmdale Silt	3	19	tr.	11	13	22	0	0	40	0	0	0	17	9	74
Ancient Mississippi Valley (fig. 3) Roxana Silt,															
Zone IV Roxana Silt,	13	18	1	12	9	26	0	2	45	1	tr.	1	17	7	76
Zone III Roxana Silt,	9	18	2	10	10	25	tr.	1	49	1	tr.	1	18	9	73
Zone II Roxana Silt,	16	16	3	11	11	19	tr.	2	51	1	tr.	1	16	10	74
Zone I	14	17	4	10	14	26	tr.	1	32	1	tr.	1	13	8	79
Roxana Silt, undiff. Ancient Iowa Valley (fig. 3)	14	19	2	8	11	28	tr.	1	45	2	tr.	1	14	8	7 8
Roxana Silt, undiff. Ohio River Valley	5	33	7	8	12	29	8	0	31	tr.	0	0	-	-	-
Roxana Silt, undiff. Wabash River Valley	3	9	10	26	8	18	1	tr.	31	0	0	0	14	7	79
Roxana Silt, undiff.	1	24	1	12	16	21	0	0	49	0	0	0	-	-	_

TABLE 3—Continued

			·	(cent anal		-		_)		ight nerals	
Depositional province and stratigraphic unit	Number of samples	Black opaque	Tourmaline	Zircon	Garnet	Epidote	Staurolite	Actinolite	Hornblende	Enstatite	Hypersthene	Diopside- augite	K-feldspar	Na-Ca feldspar	Others
Upper Mississippi Valley Late Woodford- ian outwash Illinois and Rock River Valleys	6	20	1	2	15	37	tr.	1	36	1	1	1	16	8	76
Late Woodford- dian outwash Illinois River Vall	6 ey	30	1	2	16	13	0	1	52	2	2	5	17	11	72
Early Woodford ian outwash Northwestern Illino	26	15	2	3	16	15	tr.	1	51	1	3	2	16	9	75
Winnebago outwash	3	29	1	2	19	13	0	0	53	-	-	11	21	13	66

TABLE 4 — CLAY MINERAL AND CARBONATE X-RAY ANALYSES

) ,	ı	,	i																													
ls in tion	ercent aolinite nd chlorite	1	22:	2 5	‡ C	14	8	13	16	12	56	8	7	18	18	14	14	14	14	17	13	10	ω	12	18	18	15	σ	13	7	12	23
Clay minerals in < 2 \mu fraction	ercent llite	- 1	77	0 1	~ e	2 8	17	19	16	27	52	13	14	14	24	53	R	R	32	8	23	53	46	27	73	80	81	15	22	11	12	15
Clay < 2,	ercent ontmoril- onite	u	99	107	, .č	26	52	89	89	19	49	62	65	89	58	57	61	19	61	47	80	19	4	61	6	0	4	92	65	89	9/	62
	. I. Ratio	I	4.0		٠ د د	4.	4	1.0	7.	1.6	.7	4	ı.	ů.	6.	1.4	1.2	1:1	1:1	1.4	1.4	2.0	9°8	1.5	2.9	2.4	3.5	0.1	1.1	۳.	.7	ı,
	on Dolo-	mite	م ز	, i	40	13	'	ı	•	ı	ı	ŧ	•	,	•	•	•	•	•	ı	8	•	•	•	8	99	임	1	ı	1	ı	1
action counts	< 2 μ Fraction	Calcite	1 6	۱ کا	5	} '	1	•	ı	,	1	1	ı	•	ı	1	1	ı	ı	ı		ı	ı	•	12	ı	•	•	•	ı		ı
X-ray diffraction intensity in counts per second		mite	100	267	70	75		•	•	,	ı	ı	•	ı	ı	•	•	•	ı	ı	110	•	,	•	170	82	165	1	,	ı	,	ı
X-1 inte	Bulk sample	Calcite	18	2 '	י כ	90		ı	,	,	1	•	,	•	,	1	ı		ı	ı	15	1	ı	1	15	4	23	,	ı	,	ı	•
	omper	1	311	20,4	325	326	332	333	341	345	342A	342B	342C	342D	342E	342F	3426	342H	342I	3427	351	352	329	360	361	362	366	373	379	380	381	382
ls in	ercent aolinite nd chlorite	1	77	3 5	3 5	12	19	15	8	21	72	16	16	16	17	14	14	17	19	14	8	16	15	6	14	o	14	18	13	+	7	12
Clay minerals in $< 2 \mu$ fraction	ercent llite		27	7 -	2 5	22	18	18	8	27	8	27	22	19	35	24	48	77	8	37	တ္တ	88	ଷ	18	12	7	15	55	16	+	9	11
Clay <2	ercent ontmoril- onite	0	22	5 5	15	63	63	29	9	52	44	57	62	65	51	62	88	62	55	49	20	53	65	73	74	84	7	9	71	+	87	11
	oitsA .I .	I	φ.	, (9 4	0	9	φ.		6,	1.0	1.0	6.	ထ္	1,3	1,1	2,3	ထ္	6.	8	1.0	1.0	6.	1.3	9.	9.	.7	φ	φ	+	τĵ	9.
	μ ion Dolo-	mite	15		7		15	22	140		,	•	1	9		•	1	•	14	ស			ı	•	1	1	,	•	17	+	7	15
X-ray diffraction intensity in counts per second	$< 2 \mu$ Fraction	Calcite	1	7	t %	8 00	72	7	ı	1		•		1	1	ı	•	ı	•	•	1	1	1	ı	•	ı	1	•	1	#	,	ı
X-ray diffracti intensity in cou per second		mite	9	ı ı	Ĭ.	42	35	9	13	1		ı	,	82	9	ı	ı	ı	110	125		•	•	ı	ı	i	ı	•	155	8	120	70
X-: inte	Bulk sample	Calcite	77	7 '	. 1	13		œ	45	1	ı	•	1	4	1	1		,		ı	1	1	1	ı	1		1	•	ı	•	ı	•
	ample umber	J	∞ δ	ν α α	5 6	888	2	104	11	12	17	18	19	8	7	22	23	œ	4	83	84	82	98	87	88	86	66	9	101	102	103	104

22112777748 8 9 2 2 1 1 1 2 2 1 1 1 2 2 1 1 2 2 1 1 2 2 1 1 2 2 1 1 2 2 1 1 2 2 1 1 2 2 1 2 2 1 2 2 1 2 2 1 2 2 1 2 2 1 2 2 2 2 1 2	17 19 11
82482888217	30 88 30
\$40874058 82674874171 8682685811 8688866 640874058 86874874171 868868686866	64 53 59
1411646 1 11 0 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	.8 1.0 1.7
25, 24, 27, 28, 27, 28, 28, 28, 28, 28, 28, 28, 28, 28, 28	101
17, 11, 18, 11, 11, 18, 11, 11, 11, 11, 11	1 1 1
1120 1150 1150 1330 1330 1330 1330 1330 133	10 - 240
1888888	1 1 1
3885 3885 38865 3886 3886 3886 3887 3887 3887 3887 3887	485 486 487
27.45.10.20.20.20.20.20.20.20.20.20.20.20.20.20	5°6 8°6 8°6
088888888	43 14 14
87 87 87 87 87 87 87 87 87 87 87 87 87 8	37 82 66
0100001410 100000 010000000000000000000	2.1
1181111 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	11 28 -
ווויווי ביין ביין ביין ביין ביין ביין בי	اھا
265 2000 2000 2000 2000 2000 2000 2000 2	160
	12
105 106 107 108 109 109 1116 1117 1128 1133 1134 1135 1135 1135 1135 1135 1135	287 288 310

Material sufficiently weathered that a quantitative evaluation of clay mineral composition is impractical. Dune sand containing essentially vermiculite and kaolinite in the clay minerals. No data.

ls in	chlorite chlorite	кво	24	8 :	<u>'</u> α	17	15	15	∞	12	13	6	6	15	18	18	14	_ 1	7	on o	a	10	17	18	19	=	18	15	2	16	12
Clay minerals in $< 2\mu$ fraction	tnec eti	Perc	တ္တ	24	2 5	1=	18	12	3	16	33	23	19	54	2	20	18	12	17	ର ଚ	3	36	22	33	23	17	3	55	24	8	55
Clay c	tnec tmoril- ite		46	8	8 2	52	67	73	29	69	49	89	72	. 61	75	62	89	81	9/	7,	6	54	28	26	28	72	53	63	99	28	63
	oiteA .I	. • d	ω.	٥, ١	ο α	4	ω	۰.	2.0	٠.	1:1	1.6	1.4	1:1	4	٠.	6	7.	1.6	9.	o G	3.2	0:	٥.	ထ	1:1	1:1	1.0	1.6	1.1	1.0
	د	Dolo- mite				,	ı	12	22	15	12	1	21	23	ı	ı	ı	13	55	۰:	01	•		4		12		16	77	ഹ	7
X-ray diffraction intensity in counts per second	$< 2 \mu$ Fraction	Calcite	ı			•	,	ı	ı	•	•	1	50	9	1	•	,	ω	•	1	1			•	•	•	1	ı	ı	•	
X-ray diffraction ntensity in count per second	ω	Dolo- mite	1		, ,		ı	75	82	100	92	ı	21	82	22	•	10	40	ß	8 :	ဂ္ဂ	1	1	1	•	æ	•	65	150	8	8
inte	Bulk sample	Calcite				1	•	•	ı		•	1	ω	33	D.		ı	13	•	<u>ا</u> د	n	•	•	,	,	1	,	ı		•	12
		lme2 Imun	623	624	623 625	626A	626B	626C	626D	627	628	629	633	634	635	637	636	640	642	644	040	647	652	653	654	655	929	657	6 58	629	099
ls in tion	chlorite chlorite	к 90]	12	20	4 %	3 %	12	33	0,	*	56	12	13	0	17	23	17	10	13	75	<u>ي</u>	17	11	6	23	16	19	12	11	17	*
Clay minerals in $< 2 \mu$ fraction	tnə: əti	Perc	43	38	4 4	88	88	11	16	*	17	27	23	16	12	19	14	19	52	15	У.	ω	22	12	ဓ	12	6	12	16	13	*
Clay < 2/	ent cmoril- ite		45	12	12	`=	: 89	58	75	*	57	19	64	75	89	28	69	71	62	61	20	75	29	79	47	72	72	92	73	70	*
	oitsA .l	υ • α	2.4	ឃុំ	ω. Γ	7		α.	1.2	4	4.	1.5	1.2	1.1	9.	9.	9.	1.3	1.3	4.	ů.	ლ	1.4	.7	6.	τů	ლ.	٠,	1.0	ď.	•
,	μ ion	Dolo- mite	18	10	4 წ	3 5	17	,	16	•	•	30	23	•	1	,	1	15	17	•	ı	•	33	33	1	•	,	വ	01	ı	1
X-ray diffraction intensity in counts per second	$< 2 \mu$ Fraction	Calcite	1	19	23 7	2 4	} '	ı	ı	•	•	22	ı	•	,	,	ı	,	55			,	10	56	1	•		•	,	•	1
X-ray diffracti ntensity in cou per second		Dolo- mite	85	36	9	8 5	4 6	ı	75		ı	35	9	•	•	1	•	75	4			ı	170	6	•	•		ဓ	55		ı
X- int	Bulk sample	Calcite	12	တ္တ	8 6 7	, ç	3 '		1	•	,	Ŋ	1		1	ı	ı		16			ı	22	24	•	•	•	1	9	1	ı
	ole Jer	gms2 dmun	488	489	490	493	501	504	505	511	512	513	514	515	523	524	525	526	527	533	534	535	536	537	544	545	546	547	548	553	554

ç	5 4	2	2 6	28	12	0.	101	=	7	œ	23	55	23	78	10	11	15	11	18	28	15	7	ω	13	12	12	15	13	12	12	15	9	16	13	14	15	23	19	17
ć	15	6	20	8	24	22	8	8	32	56	18	14	ស	14	12	16	8	71	77	21	3	17	18	8	40	88	82	87	88	77	36	99	65	35	88	35	54	99	70
6	2.5	65	62	26	64	65	62	55	99	99	29	49	72	28	78	73	51	89	61	53	62	92	74	49	48		1	٠	1	11	49	84	19	55	28	53	23	15	13
~		ω,		1.0	1.3	1.6	2.0	2.0	2.2	2,3	r.	က္	7	ლ	ω.	6.	1.6	1.3	œ	ທີ	1.1	1.5	1.6	2.0	2.2	4.7	8	4.3	5.1	4.3	1.6	1.6	8	1.5	1,3	1,4	1.6	2.3	2.7
١	15	00	17	11	12	10	11	11	•	ı	f	,	•	ı	1	•	9	1	•	•	7	01	10	1	15	32	42	50	1	17	14	21	55	16	37	56	13	35	15
•	•	ì	,	1	1	1		1	•		,	•		ı	ı	1		1	•		70	1	ı	ı	ı	7	19	10	12	7	ı	ω	11	•	10	ω	13	15	7
ı	55	70	3	45	115	80	06	8	•	1	•	•	ı	1	ı	1	115	,	•	•	82	65	75	ı	100	105	20	65	380	410	20	95	310	75	9	75	70	140	180
,	ı	1	ഹ	ı	•	•	•	•	1	,	•	•	•	17	•	•	ı		1	1	22	1	•	•	•	12	9	8	17	17	12	14	10	43	13	16	10	16	95
199	663	664	999	999	299	899	699	929	671	672	692	693	694	695	719	720	723	761	776	790	801	802	803	804	802	828	824	822	827	829	830	831	832	833	834	835	836	845	846
40	27	22	8	27	27	23	ဓ	16	14	12	18	23	88	16	27	23	14	18	16	18	19	31	73	88	13	23	19	16	16	6	18	6	77	18	50	77	16	ω	- ∞
. %	42	26	10	15	34	54	17	14	6	6	15	33	36	19	ß	8	14	27	56	13	17	23	23	50	15	22	56	24	23	37	57	55	31	24	53	35	27	ო	12
50	31	55	64	28	33	53	53	70	77	79	4	42	98	65	48	42	72	22	28	69	64	46	54	25	72	22	55	9	61	54	22	69	48	28	51	47	57	82	80
α	1:1	1.8	ლ.	4.	o.	.7	4.	9.	4.	ů.	۰.	1.0	φ	ထ္	9.	1.0	٠.	1.0	1.1	τů	9.	ι.	.7	ι.	ထ္	٠,	6.	1.0	1.0	2.6	2.1	1.6	1.0	٥.	1.0	1.0	1.1	•5	1.0
٠	12	•	4	1	•	1	ı	i	•	1	Í	27	•	•	•	32	1	,	16	•	ı	ŀ	20	52	12	•	11	14	10	10	55	,	•	•	ı	•	6	1	14
,	,	•	ı	ı	,	,		1	r	1	•	•	•	1	,	1	,	,	•	,	ı	,	1	ı	•	ı	1	ı	,	,	•	1	ı	•	,	ı	ı		ı
20	135	1	1	80	96	•	1	ı	•	•	1	140	•	ı	•	130		•	115	•			9	45	30	,	24	65	20	120	145	•		ı			30	1	110
ı	1		,	15				,	ı	•	1			•	ţ		•			•	ı	1		0	•	1		12	1	1	•	ı	ı	1		1	ı		
556	557	561	562	564	265	267	268	269	570	571	572	573	574	575	216	211	578	579	580	587	588	589	591	592	593	594	595	296	297	298	299	009	909	607	809	609	610	617	618

Material sufficiently weathered that a quantitative evaluation of clay mineral composition is impractical. No data. Organic-rich silts composed essentially of vermiculite and kaolinite in the clay minerals.

inued
귣
ຮື
4
ų
TABI

4	s in ion	rcent slinite d chlorite	ks	14	77	£ :	2 5	<u>.</u>	3 5	2 8	13	13	15	17	55	18	50	77	15	9	18	16	13	23	13	10	0	ខ	8	10	
	Clay minerals in $< 2\mu$ fraction	rcent lite		11	σ,	15	Ξ:	2 1	~ <u>u</u>	, ,	33	9	33	£3	55	27	48	12	4	5	53	ß	27	24	ဓ	37	8	4	9	33	
	Clay m	rcent ntmoril- ntte	IOM	75	79	۲ ا	=	- 6	e 6	7 6	3	57	20	28	26	55	35	4	7	25	23	26	9	53	22	23	22	4	36	49	
		I. Ratio	ρ.	٦.	ທີ	۰,	ه ۱	ດຸເ	• ·		1.1	1.5	1.5	6.	.7	1.0	1.6	ທີ	•	•	1.1	1.1	1.4	.7	1.5	2.4	2.7	5. 8	9.	1.4	
		n Dolo-	mite		٠	•	•			•	13	12	16	14	c o	15	7	1	•	œ	ß	78	16	16	7	10	•	16	•	•	
	X-ray diffraction intensity in counts per second	$< 2 \mu$ Fraction	Calcite	ı	1	•	t	1			ı	1		•	ı	•	,	1	1			•		•	ı	•		ı	•		
	ray di ensity per se	6 Dolo-	mite	12	24	ı				ç	35	100	65	195	2	32	110	•	•	15	120	160	100	120	5	105	160	155	1	1	
0	X- int	Bulk sample	Calcite	1	•	,			• 1	1	,	ı						ı	•				•	•	1			1	ı	,	
ontinu		mple mper	ınu	616	920	921	922	5 73 6 73 7 7	4 2 4	926	927	928	929	930	931	932	933	934	932	936	937	938	626	940	941	942	943	944	955	926	
TABLE 4 — Continued	ls in	rcent olinite d chlorite	ķЯ	21	19	18	Ξ:	: P	16	7 5	19	27	22	52	27	56	18	14	24	22	18	16	13	13	16	19	19	50	19	8	
IAB	Clay minerals in $< 2\mu$ fraction	rcent lite		20	88	2	8 9	5 6	2 K	3 5	1 %	14	19	16	17	18	17	2	16	14	7	21	8	18	19	55	77	55	22	24	
	Clay m < 2 \mu	rcent ntmoril- nite	ЮШ	59	53	89	69 \	8 9	2 2	3 2	5 6	56	26	62	26	26	65	92	9	4	61	63	29	69	65	26	62	28	29	26	
		I. Ratio	•a	.7	1.0	1:1	1.2		· •	o a	0	4.	ι	ů.	4.	4.	ď.	ď.	4.	4	φ.	6.	1.1	ထ္	ထ္	φ	ထ္	.7	ထ္	Φ.	
		tion Dolo-	mite	5	10	9	2	2;	7.	ָ ק	28	'	ı	ı	•	ı	ı	•	1	•	12	11	12	11	8	14	10	12	11	2	
	X-ray diffraction intensity in counts per second	< 2 μ Fraction	Calcite	•	•	ı	1	•			,	ı	•	ı	1	•	1	,	1	ı		ı		ı	•	ı	•	,	•	1	
	X-ray diffracti ntensity in cou per second	e Dolo-	mite	36	80	8	120	ဥ္က ဗ	S 4	S K	5.4	· •	1	•	•	•	•	•	1	•	8	75	115	140	105	8	8	80	8	20	
	x- int	Bulk sample	Calcite	•		24						•	,	,	1	•	10	11	10		•	ı	1	1	•	•	•		ı	9	
	 	wper wble	mu	347	848	349	820	851	222	25.0	855	857	858	829	860	861	862	863	364	865	866	867	868	698	870	871	872	873	874	875	

8 / 8 0 1 1 1 1 2 2 8 2 1 4 2 1 1 2 2 7 7 7 7 7 7 7 7 7 7 7 7 7 7 7
888 9 11 11 12 12 12 12 12 13 13 13 13 14 15 15 15 15 15 15 15 15 15 15 15 15 15
22488665556 868848511024 46288888891186 855515588817
40 6 6 7 4 4 6 6 6 8 6 6 6 6 7 6 7 6 6 6 6 6 6 7 6 7
01
- 01 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1 - 1
5 20 1
188 1 1 1 188 18 18 18 18 18 18 18 18 18
959 960 975 975 982 983 984 987 988 989 989 989 989 989 989 989 989
84411 911814 91201 81211 91211 <t< td=""></t<>
4 2 4 4 4 2 8 8 8 4 2 4 8 4 8 4 8 4 8 4
8444474646 77 - 42 * 8786 87 - 87446 7446 7446 87 87 87 87 87 87 87 87 87 87 87 87 87
00000000000000000000000000000000000000
1222 1111 1 2 2 2 2 3 3 2 3 3 2 3 3 2 3 3 3 3
1 1 1 1 1 1 1 1 1 1 2 2 2 4 2 2 2 3 1 1 1 1 2 1 1 1 1 1 1 1 1 1 1 1
165 105 105 105 105 105 105 105 105 105 10
01 1 1 1 1 1 1 1 1 1
878 879 881 882 883 884 884 887 887 889 890 890 890 890 890 890 890 890 890

* Material sufficiently weathered that a quantitative evaluation of clay mineral composition is impractical.

TABLE 4 — Continued

			ı																													
ls in	tne inite etiroldo		18	54	46	23	23	7	17	12	10	12	17	12	14	14	ω	56	18	46	16	14	58	17	33	12	14	17	10	17	12	33
Clay minerals in $< 2 \mu$ fraction		Perce	64	33	2	27	18	40	62	72	62	61	26	61	62	53	31	35	32	47	26	62	32	22	33	54	56	18	17	71	32	30
Clay < 2	-liron	Perce nontr	18	13	34	20	29	ဗ္ဗ	77	13	78	27	27	27	54	33	61	42	47	7	88	24	40	88	34	64	09	65	73	62	20	37
	oits8 .		2.4	4.	ო.	ထ္	r.	1.4	2.5	3,3	4.1	တ္	2.2	3.4	2.9	5.6	2.5	٥.	1.3	٠.	2,3	2.9	œ.	2.2	.7	1.4	1.2	.7	1.2	φ.	1.6	9•
	g	Dol∵- mite	•	,	1	1			8				22					ı	ı			13	,	ı	•		ı	1	ı	1	1	
X-ray diffraction intensity in counts per second	$< 2 \mu$ Fraction	Calcite			•			ι	ı	ı		1	1	ı	•	•	77	ı	•		1	9		ı	ı	1	1	,	ı			
ray diffra ensity in per second	a)	Jolo- nite	,	•	•	•	27	80	75	135	8	125	100	80	120	120	145			,	15	82	•		,		1	ı		ı		
X-int	Bulk samole	Calcite	1	,		1	•	•		ω	8	35	12	52	15	15	35		1			•	•	ι	•		ı			1		1
		dme2	1132	1135	1136	1137	1138	1139	1140	1141	1142	1143	1144	1145	1146	1147	1148	1149	1150	1154	1155	1156	1159	1160	1168	1171	1173	1176	1178	1182	1183	1184
s in ion	tne inite chlorite		6	10	12	16	6	ω	ເດ	ω	10	10	ω	33	28	Ξ	35	29	23	8	18	55	12	15	6	25	15	33	13	35	31	56
Clay minerals in <2 \mu fraction		Perce	2	12	18	22	21	23	22	19	54	23	22	37	27	8	56	24	17	10	13	22	18	5	21	33	31	56	21	45	54	တ္တ
Clay m	-liton	Perce mont	81	78	2	62	70	69	7	73	99	29	20	ဓ	45	55	45	47	9	2	69	26	70	26	2	40	54	36	99	55	45	44
	oiteA .	D° I	۲.	φ	6	1.0	1.5	1.9	3,3	1.5	1.6	1.5	1.8	ω.	.7	2.1	•	9	r.	ღ	ຜ	.7	1.0	1,3	1,5	٥.	1.4	τ.	1.1	٥.	ທ໌	ထ္
X-ray diffraction intensity in counts per second	$< 2 \mu$	Dolo- mite	14	9	ı	2	20	21	14	•	99	27	32	1		1	•	ı	ı	,	ı	1	•	•	•	,	•	1		1	1	
	> Pragr	Calcite		,	•		ı	,		ı	•		1		1			ı	1	1		ı	٠		ı	,		1	ı		•	
ray dif ensity per sec	·	Dolo- mite	9	55	17	80	110	130	105	52	160	40	47	,	•	•	•	•	•	,	•	,	1	•	ı	•	•	1	1		•	ı
X- int	Bulk	Calcite			•		•	ı	ı	13	•		ı	ı	1	ı		•	ı	ı				1	1	ı	ı	•		1	ı	1
	1e Te	dms2 dmun	1016	1017	1018	1019	1020	1021	1022	1023	1024	1025	1026	1028	1029	1030	1048	1049	1050	1051	1052	1053	1054	1055	1056	1059	1060	101	1072	1078	1079	1080

10 23 10 12 12 13 13	98 * 911111299	21 111 120 7 8 8 8 4 8 4 8 4 8 4 8 8 4 8 8 8 8 8 8	20 11 11 11 10 9 9 11 11 11 11 11 11 11 11 11 11 11 11
111 113 116 110 111 111 115	118 119 119 119 119 119 119	26 100 100 100 100 100 100 100 100 100 10	11 13 10 11 11 11
668 688 688 633	24 + 17 + 14 + 14 + 14 + 14 + 14 + 14 + 1	47.007.00 400.00 400.00 400.00 400.00 400.00	69 75 71 71 75 80 80 75
1	8048407774	122 138226	4.7.2.1.7.8.8.8
15	1 1 1 2 2 2 2 2 2 2 1 1 1 1 2 2 2 2 2 2	11011	22 21 21 16 18 12
	27 27 25 25 25 25 25 25 25 25 25 25 25 25 25		
% 488 88	35 55 57 57 57 57 57 50 50 50 50 50 50 50 50 50 50 50 50 50	11111 60031	55 110 105 140 125 40
	122 7 1 1 1 1 2 2 7 1 1 1		1 10 10 10 10 10 10 10 10 10 10 10 10 10
1185 1191 1192 1193 1194 1196 1198 1214	1216 1217 1218 1220 1220 1221 1223 1223	1229 1230 1231 1232 1233 1234 1235 1236 1239	1241 1242 1243 1244 1245 1246 1248
23 23 23 18 16 16 15 15	19 118 12 14 14 15	118 119 122 122 138 143 143	81 2 3 3 3 3 3 4 5 8 5 8 5 8 5 8 5 8 5 8 5 8 5 8 5 8 5
53 63 63 71 70 74 63	88 * 05	77 75 70 70 70 70 72 72 78	77 72 72 73 74 71 71
22 22 24 25 25 26 27 27 27 27 27 27 27 27 27 27 27 27 27	23 * 25 * 25 * 25 * 25 * 25 * 25 * 25 *	31 31 11 11 11 11 11 11	22. 7. 25. 6 9 8 5 9 8 5 9 8 9 9 9 9 9 9 9 9 9 9 9 9
	41. 41. 42. 68. 68. 68. 68. 68. 68. 68. 68. 68. 68	444444444 844444444	8.00 6.00 7.00 8.00 8.00 8.00 8.00 8.00 8.00 8
122880	152 111 20 20	2	
55 105 105 60 60 60 75 100	55 50 110	8.11111111	
5555	10		

Material sufficiently weathered that a quantitative evaluation of clay mineral composition is impractical. No data.

TABLE 5—AVERAGES OF CLAY MINERAL DATA BY STRATIGRAPHIC UNIT AND DEPOSITIONAL PROVINCE

	% i	n clay min	neral	
Stratigraphic unit and depositional province	Mont- morill- onite	Illite	Kaolin- ite + chlorite	D.I. ratio
Richland Loess and Upper Zone Peoria Loess	1911/1920			
Illinois River Valley	42	41	17	3.0
Mississippi Valley above Alton	80	11	9	.7
Mississippi Valley below Alton	66	26	8	2.3
Peoria Loess (major part) and Morton Loess				
Illinois River Valley	54	30	16	1.3
Mississippi Valley above Alton	70	19	11	1.2
Mississippi Valley below Alton	71	22	7	1.5
Wabash River Valley	26	61	13	3.1
Ohio River Valley	9	71	20	2.5
Farmdale Silt				
Illinois River Valley	62	17	21	•5
Transition Zone				
Ohio and Wabash Valleys	21	55	24	1.4
Roxana Silt				
Ancient Mississippi Valley (fig. 3)	70	14	16	•6
Ancient Iowa Valley (fig. 3)	64	16	20	•5
Wabash River Valley	55	22	23	.6
Ohio River Valley	30	38	32	.8

TABLE 6 — CHEMICAL ANALYSES OF ILLINOIS LOESSES

Sample Number	sio ₂	Tio2	A1 ₂ 0 ₃	Fe ₂ 0 ₃	FeO	MgO	CaO	$^{\rm Na}_2{}^{\rm O}$	K ₂ 0	H ₂ 0+ (>110°)	H ₂ 0- (<110°)	c ₀ 2	so ₃	Ignition loss	Total
PEOR IA	LOESS														
11	64.99	9.	90.6	2.09	.58	4.05	68.9	1.66	2.01	1.33	.67	7.47	Ξ.	8.70	100,63
19	75.60	• 78	11.62	3,41	34	.85	1.08	1.67	2.41	2.79	1.41	8	ි.	2,63	100,39
108	58,88	• 52	8.42	1,65	.68	5.73	8.98	1.59	1.58	1.10	.42	11.28	.1	12,38	100.44
342-F	74.62	.81	12,10	3.94	44.	1.07	.95	1.29	2,58	3.24	1.91	8	8	3.07	100.90
470	58.58	.48	8.55	1.81	.54	5.87	8.40	1.36	1,65	1.78	• 79	11.36	ş	12.79	100,02
473	69.53	.73	10.87	2.96	.28	2.73	3,75	1,55	1.82	2,77	1,36	3.87	8	6.23	100.45
536	64.90	99.	8.71	1.85	.61	3.79	7.16	1.53	1.64	1.65	.64	8.15	.01	9.40	100,25
537	62,95	48	9.21	2.11	5.	3.76	7.96	1.64	1,52	1.95	.82	8.57	8	10.14	100,31
265	63.48	69.	8.76	2.16	•62	4.53	6.81	1.54	1.88	1.81	• 58	8.47	• 05	9.06	100,43
266	54.67	.54	6. 90	2.06	•62	6.57	10.46	1.17	1.75	1.75	•39	13.99	• 02	15.44	100,18
MORTON	LOESS														
136	55,23	• 58	7.86		• 78	6.51	10.00	1.63	1.73	1.84	.43	12.99	60.	14.46	100.51
137	57.10	86•	7,75	1.27	1.20	5.88	8.83	1.31	1.83	2,50	•62	11,30	.52	13.75	99.90
FARMDALE SILI	LE SILT														
691	73.76	69.	13,14	3.52	99.	1.00	.81	83	1.88	4.31	2.60	8	8	4.29	100,58
695	76.04	•86	11,15	1.78	1.07	86.	1,55	1.81	2,52	2.63	%	8	8	2.57	100,33
ROXANA	SILT														
∞	70.87	.63	89.6	2.26	.51	2,57	4.31	1.81	2.21	1.29	.67	4.03	60.	5.38	100,23
8AA	75.09	69.	11,55	2.57	.61	1.05	2.05	1.76	2.21	2.36	1.30	•65	•03	2.67	100.25
10	63.91	.61	8,60	2.03	.56	3.94	7.38	1.40	2. 04	1.52	.63	8.37	9	9,65	100,12
12	75,12	• 78	10,06	2,38	.71	1.62	2.58	1.74	2.40	1.59	.74	1.55	8.	2,95	100,34
18	76.73	85	11.00	3.13	.38	.77	1.00	1.60	2,35	2.66	1.11	8	\$	2.48	100.29
66	75.80	.64	11.79	2,70	• 58	1.17	1.43	1.76	2.04	2.13	1.54	.10	.11	2.39	100,30
101	65.65	99.	9.72	2,95	.42	3.82	5.49	1.59	5. 06	1.50	68	6.35	.13	8.03	100,39
103	64.97	• 26	9.52	2.59	.42	4.24	5.99	1.49	1.77	1.68	1.22	7.26	.11	8,95	100,50
105	76,65	.91	10.89	2.59	.81	1.00	1.59	1.91	2.62	1.47	• 78	ş	.12	1.63	100,60
342-D	76.44	• 76	11.09	3.16	.41	.82	1.06	1.88	2,31	2.59	1.61	8	8.	2,55	100,48
462	75,90	.68	11,37	2,75	.53	1.08	1.29	1.78	2.18	2.92	1.42	8	8	2.81	100,37
463	73,35	.71	10.69	2,50	• 52	1.07	3.48	1.85	2.25	2.71	1.10	1.83	8	4.03	100.42
466	68.39	52	9.54	2.27	.52	3.25	4.91	1.61	2.04	2,25	• 79	5.4	8	7.36	100.44
467	77.07	• 78	10.90	2.69	•26	1.07	1.37	1.75	2.26	2,14	1.44	8	•05	1.97	100.42

SELECTED GEOLOGIC SECTIONS

Following are fifteen of the measured geologic sections used in this report; fifteen additional sections also used in this study were published in Illinois State Geological Survey Circulars 285 and 304 in 1960. The numbers enclosed in parentheses, for example (P-565), are sample numbers used in the tables and illustrations in this report. The sections are arranged alphabetically by name.

DALLAS CITY SECTION	Thickness
Measured in overburden of quarry in SW4 NE4 SW4	(feet)
sec. 36, T. 8 N., R. 7 W., Henderson County,	Pleistocene Series
Illinois, 1961.	Wisconsinan Stage
Thickness	Woodfordian Substage
(feet)	Bloomington Moraine
Pleistocene Series	10. Till, pink, bouldery, calcareous,
Wisconsinan Stage	compact 10.0
Woodfordian Substage	9. Gravel, tan, calcareous, locally
Peoria Loess	cemented in upper 1 to 2 feet,
7. Loess, coarse, friable, calcareous,	generally fine grained but with
yellow-tan, massive but with indis-	cobbles up to 3 to 4 inches in
tinct bedding in lower part, sparse-	diameter 8.0
ly fossiliferous; surface soil in	Shelbyville Moraine
top (P-1248 middle; P-1247, 5 ft.	8. Till, gray to gray-tan, bouldery,
above base; P-1246 base) 22.0	massive, calcareous (P-558); 2-foot oxidized zone at top 20.0
Loess, coarse, calcareous, massive,	2-foot oxidized zone at top 20.0 Morton Loess
gray to light tan, brown mottled;	7. Moss, clay and silt, lenticular,
abundant limonite tubules; a few	light gray to tan, calcareous;
fragmentary snail shells (P-1245	organic matter throughout, but
top; P-1244; P-1243 base) 8.0	for a few millimeters the moss
Silt, gray, mottled brown, very	is a compact mat (W-483,
weakly calcareous; contains limon-	20,500 ± 600 years B.P.) 0.1
ite tubules; limonite streak at	6. Loess, dove gray streaked with
top (P-1242) 1.5	tan, massive, slowly calcareous
Altonian Substage	to vigorously calcareous in upper
Roxana Silt	3 inches (P-557), fossiliferous 2.5
4. Clayey silt, compact, massive,	Farmdalian Substage
gray, noncalcareous; limonite	Farmdale peat and silt
streak at top and gradational at	5. Peat and silt; silty black peat
base (P-1241) 1.5	at base, streaked with brown and
Silt with some sand, becoming sandy in lower part, noncalcareous,	tan silt in upper part and
brown in lower part, noncarcareous,	chocolate-brown silt at top (W-406,
brown mottled at top, massive, com-	25,150 ± 700 years B.P.) (P-555;
pact; contains some clay (P-1240	P-556) 3.5
top; P-1239 base) 2.5	 Silt, medium gray to purple-tan,
Illinoian Stage	massive, compact (P-554) 2.0
2. Sangamon Soil developed in till,	Altonian Substage
leached, brown; Sangamon Soil in	Roxana Silt
top; B-zone (P-1238) strongly de-	Silt, clayey silt and silty clay,
veloped with splotches of brown	very dark gray and finely mottled;
clay in gray matrix, grading down-	pebbles very rare; organic staining
ward into leached, brown till 6.0	throughout; platy to crumb structure;
1. Till, silty, calcareous, blocky,	noncalcareous, gradational contacts
tan, brown, and gray; contains few	(P-553) 1.5
cobbles (P-1237); rests on Salem	Illinoian Stage
Limestone 5.0	Colluvium of till and silt grading upward into accretion gley; blue-
Total 46.5	gray to green-gray, grading upward
10101 40.0	to dark gray to black; becomes
DANVERS SECTION	quite clayey in top with small
Measured in NW1 NE1 NW1 sec. 32, T. 25 N., R.	blocky structure, noncalcareous
1 W., Woodford County, Illinois, in cutbank on	(P-551 lower; P-552 upper) Sangamon Soil in top 5.0
south side of Rock Creek (near the locality of	Sangamon Soil in top 5.0
section 22 of Horberg, 1953), 1959.	

			51
· -	kness	FRENCH VILLAGE SECTION	
1. Till, calcareous, becoming weakly calcareous at top, irregularly oxidized in upper part; gray pen-	eet)	Measured in large pit in Mississippi Valle in NE ¹ 4 NW ¹ 4 NW ¹ 4 sec. 25, T. 2 N., R. 9 W., County, Illinois, 1961.	y bluff St.Clair
dants extend downward into tan (P-550)	4.0		kness eet)
• •	56.6	Pleistocene Series	
Total	30.0	Wisconsinan Stage	
DEPUE SECTION		Woodfordian Substage Peoria Loess	
Measured in SW¼ SE¼ SE¼ sec. 27, T. 16 N., 10 E., Bureau County, Illinois, 1959.	R.	9. Loess, gray-tan to yellow-tan with pinkish cast in lower part, calcareous, fossiliferous, indistinct	
	kness eet)	textural bands in lower part (samples range upward from P-1022	
Pleistocene Series		at base, to P-1026 at 6 feet below	00 0
Wisconsinan Stage		top)	23.0
Woodfordian Substage		 Loess, pale pinkish tan, massive, calcareous, sparsely fossiliferous, 	
15. Silt (Lake Illinois silts), gray,	4 0	gradational contact at base (P-1020	
calcareous, thin bedded (P-385)	4.0	lower; P-1021 upper)	2.5
Bloomington Moraine 14. Till, pink, compact, calcareous		Altonian Substage	
(P-384)	4.0	Roxana Loess	
Morton Loess		Zone IV	
13. Sand, very fine, and silt, cal-			
careous, gray; contains fossil		7. Loess, pink to light pinkish brown,	
snails (Circular 304)	0.5	massive, compact, calcareous in lower part, fossiliferous (P-1017	
12. Loess, gray, calcareous, compact	3 5	base; P-1018 middle; P-1019 top)	5.0
(P-383)	1.5	6. Loess, gray-tan to pale pinkish	
Altonian Substage		gray-tan, massive, calcareous,	
Roxana Silt		fossiliferous, gradational top and	
Zone IV		bottom (P-1016 lower)	5.0
 Loess, weakly calcareous, purple, compact, massive (P-382) 	1.2	 Loess, pink, massive, calcareous, fossiliferous, compact (P-1015 lower) 	4.0
Zone III		Zone III	
10. Loess, gray, weakly calcareous			
locally but generally non- calcareous; locally streaks of	4.0	 Loess, coarse, containing some very fine sand, gray-tan, massive, cal- careous, fossiliferous, friable; 	
yellow-tan (P-381)	4.0	gradational contacts at top and	
Zone II		bottom (samples range from P-1010	
9. Loess, purple with rusty streaks,		at base to $P-1014$ at top)	7.0
noncalcareous, compact (P-380)	1.0	Zone II	
Zone I		2 Topog dawle pink in hang becoming	
		Loess, dark pink in base becoming lighter pink upward, massive; some	
8. Loess, gray to dark gray, compact,		clay and very fine sand; fossilifer	-
noncalcareous (P-379)	8.0	ous in upper part with sparse shells	
7. Colluvium of silt with some sand and pebbles, noncalcareous, gray	0.5	in lower part (samples range from	
Illinoian Stage	0.0	P-1004 near base to P-1009 near	
6. Sangamon Soil, B-zone, reddish		top)	12.0
brown with well developed struc-		Zone I	
ture and clay skins (P-378)	1.5		
5. Till, gray, leached, blocky		Silt, with some clay and fine sand, brown to gray-brown, massive, weakl	17
(P-377)	2.5	calcareous but contains nodules of	7
4. Till, calcareous, gray, compact,	7.0	CaCO ₂ (samples range from P-999 at	
pebbly, blocky (P-376)	7.0	base to P-1003 near top)	7.0
 Sand and gravel with limonite streaks, gray and brown (P-375) 	6.0	1. Colluvium in hole below floor of	
screaks, gray and brown (r-5/5)	0.0	nit: cilt and clay brown non-	

45.5

Till, calcareous, gray, compact, pebbly, blocky (P-374)
 Covered interval to bottom of ditch 6.0

Total

pit; silt, and clay, brown, non-calcareous (P-999-A)

Total

3.0

68.5

FULTON QUARRY SECTION

Measured in overburden of inactive quarry in NE $\frac{1}{4}$ NE $\frac{1}{4}$ SW $\frac{1}{4}$ sec. 21, T. 22 N., R. 3 E., Whiteside County, Illinois; section modified in 1959 from section measured by Paul R. Shaffer in 1954.

Thickness (feet)

2.5

6.5

Pleistocene Series Wisconsinan Stage Woodfordian Substage

Peoria Loess

8. Surface soil A- and B-zones developed in loess

 Loess, gray to light yellow-tan, coarse, mostly weakly calcareous but contains a few shell fragments (P-537 upper)

6. Loess, partly covered by slump, coarse, massive, calcareous, gray to light yellow-tan, very sparsely fossiliferous (P-536 near base) 26.0

Altonian Substage Roxana Silt

Zone I

5. Silt, coarse, pink-tan, noncalcareous, massive (P-535)

4. Silt, with some very fine sand; humic staining gradational downward suggesting an A-C soil profile, dark gray mottled with tan and some granular structure, noncalcareous (P-534) 1.

 Silt and fine sand, dark gray to dark brownish gray, friable, massive, noncalcareous(P-533 upper; P-532 lower) 3.0

Illinoian Stage

 Sangamon Soil, A-zone; sand, with some silt and clay, dark gray to graybrown; crumb structure becoming prismatic downward (P-531)

 Sand; B-zone at top (P-530), dark red to red-brown, clay-rich; grades downward into loose, tan, medium, noncalcareous sand (P-529) with locally thin gravel at base on Silurian dolomite 8.0

Total 48.7

12.0

HILLVIEW SECTION

Measured in roadcut in SW4 SW4 NE4 sec. 27, T. 12 N., R. 13 W., Greene County, Illinois, 1961.

Pleistocene Series Wisconsinan Stage Woodfordian Substage Peoria Loess

7. Loess, massive, yellow-tan to graytan with mottles and streaks of rusty brown below a 3-foot surface soil profile, calcareous; basal contact transitional through 3-foot zone (P-927 base, P-928, P-929, P-930, P-931, P-932, P-933 9 feet above base)

Altonian Substage

Roxana Silt

Zone IV

 Silt, massive, pink to pink-tan, noncalcareous but with secondary CaCO₃ nodules (P-923 base, P-924, P-925, P-926 top)

Zone III

 Silt, massive, gray-tan to dark gray, weakly calcareous; gradational contacts (P-919 base, P-920, P-921, P-922 top)

Zone II

 Silt, massive, clayey, dark pink to reddish tan, somewhat lighter in uppermost part; gradational contacts, noncalcareous to very weakly calcareous (P-916 base, P-917, P-918 top) 3.5

Zone I

 Silt and very fine sand, massive, dark gray-brown, noncalcareous (P-914 base, P-915 top)

 Colluvium of silt with some sand and clay and sparse small pebbles; gradational at top (P-913)
 2.0

Illinoian Stage

In Sangamon Soil, truncated, developed in till; sandy, clayey silt with a few pebbles and cobbles, a pebble concentrate at top; red-brown, non-calcareous (P-912)

Total

36.0

Thickness (feet)

3.5

7.0

MARION NW SECTION

Measured in cuts on Interstate 57, in cen. SW\\\ sec. 11, T. 9 S., R. 2 E., Williamson County, Illinois, 1961.

Pleistocene Series Wisconsinan Stage Woodfordian Substage Peoria Loess

5. Loess, light tan-brown, massive, leached; surface soil in top (P-1072)8.0

Altonian Substage Roxana Silt

> Loess, reddish brown with micromottling of gray-tan and black, clayey, massive, leached, gradational at top (P-1071)

Illinoian Stage

3. Till, clayey and sandy with a few dispersed pebbles, tan to brown, mottled, Mn-Fe staining on joint surfaces, leached, massive (P-1070). Sangamon Soil in top 2.5 feet; strongly developed red to red-brown B-zone 8

Pennsylvanian System

2. Yarmouth Soil developed in Pennsylvanian shale; B-zone 1 to 1.5 feet thick, brown, yellow, and black mottled with Mn-Fe pellets, massive to irregularly platy, clayey, tough, grading downward into green weathered shale that rests on nodular, weathered limestone(P-1068 green shale; P-1069 B-zone)

4.5

1.5

8.0

1.0

Thickness (feet)

 Pennsylvanian shale, limestone, and siltstone to level of interstate highway grade

Total

 $\frac{5.0}{26.5}$

6.0

NEW CITY SECTION

Measured in roadcuts in SE4 NE4 SE4 sec. 4, T. 14 N., R. 4 W., Sangamon County, Illinois, 1961.

Pleistocene Series Wisconsinan Stage Woodfordian Substage

Peoria Loess

5. Loess, tan, massive, noncalcareous; surface soil in top (P-342-E base; 342-F; 342; 342-G; 342-H; 342-I; 342-J at 4½ feet above base in B-zone)

Altonian Substage Roxana Silt

4. Loess, purple-brown, massive, clayey, noncalcareous; colluvium of brown silt with some sand in basal ½-foot (342-A base; 342-B; 341; 342-C; 342-D top) 3.0 Illinoian Stage

 Sangamon Soil, A-zone of silt with very few pebbles (P-340)
 1.5

- Accretion gley (BG-zone) at top, clayey, well structured, Mn-Fe pellets, gray above grading downward into dark reddish brown (B-zone)with gray mottling (P-339); lower part of B-zone developed in till, greenish gray, yellow-brown, and brown, sandy with some pebbles (P-338); grading downward into oxidized and leached till (P-337; P-336)
- Till, oxidized in upper part (P-335), calcareous throughout; grading downward into compact, tough, blue-gray till containing a prominent zone of sand and gravel
 9.0

Total 25.0

NEW HARMONY SECTION

Measured in pits 1 mile south of intersection in New Harmony, Posey County, Indiana (SE½ NW¼ SW¼ sec. 1, T. 5 S., R. 14 W.), 1961.

Pleistocene Series Wisconsinan Stage Woodfordian Substage

Peoria Loess

4. Loess, yellowish tan, massive, coarse, compact, friable, calcareous to within 1 foot of top, fossiliferous throughout and abundantly so in lower part; 5 feet below top an incipient soil but not leached of carbonate; concretions of CaCO₃ below the incipient soil and more sparsely throughout (samples range from P-1139 at base to P-1148 at 2 feet below top)

Altonian Substage Roxana Silt

Zones II to IV

Loess, brown, noncalcareous, massive, darker in upper part, some limonite tubules in lower part; contains very sparse Allogona profunda (P-1136 base; P-1137 middle; P-1138 top)

Zone I

 Colluvium of silt, sand, and a few pebbles, mottled brown, red, and gray, gradational upward into loess; the colluvium appears to truncate the B-zone of the Sangamon Soil below but may contain some A-zone material related to the Sangamon Soil (P-1135)

Illinoian Stage

Till, clayey, tan-gray in base grading upward into the red-brown B-zone of the truncated Sangamon Soil at top (P-1133 base; P-1134 B-zone)

Total 43.0

NORTH QUINCY SECTION

Measured in roadcuts in NW₂ NE₄ NW₄ sec. 23, T. 1 S., R. 9 W., Adams County, Illinois, 1961.

Pleistocene Series Wisconsinan Stage Woodfordian Substage Peoria Loess

10. Loess, yellow-tan, coarse, calcareous below surface soil, massive, friable, sparsely fossiliferous (P-1223 lower third; P-1222 base) 16.0

 Loess, yellow-tan and gray in indistinct beds, calcareous, sparsely fossiliferous, coarse, friable (P-1221)
 4.0

 Loess, yellow-tan, calcareous, coarse, massive, friable, very sparsely fossiliferous (P-1220 upper; P-1219 lower)

 Ioess, mottled gray and tan with some brown spots, calcareous, massive, coarse, friable (P-1218)
 4.0

 Loess, somewhat clayey, gray, dolomitic, massive; sharp contact at top but gradational at base (P-1217 top; P-1216 middle; P-1215 base)

Altonian Substage

Roxana Silt

 Silt with some clay and fine sand, purplish brown, noncalcareous, some brown mottling (P-1214)

 Sand, medium to coarse, tan-brown with slight pink cast, noncalcareous, massive; contains some angular chert fragments (P-1213)
 3.0

(P-566 near base)

8. Loess, tan to dove gray, calcareous,

compact and platy; contains shell

fragments of Succinea gelida

Morton Loess

(P-565)

20.0

3.0

Sangamon Soil developed in colluvium of clay, silt, and some sand with

red to brownish red, leached

(P-1196)

scattered pebbles of angular chert,

4.0

Thickness (feet)

3.0

5.0

0.5

0.8

1.5

Kansan Stage

1. Yarmouth Soil developed in rubble of angular chert fragments, clay and silt, dark red, tough, blocky, with Mn-Fe staining on joint surfaces; grades downward into brown clay and silt with angular chert; extends to bottom of roadcut (P-1195)3.5

Total 26.5

VARNA SECTION

Measured in NW4 NW4 NE4 sec. 16, T. 30 N., R. 1 .W., Marshall County, Illinois (cutbank exposure along tributary to Sandy Creek; ½-mile downstream from Horberg's sec. 16), by John C. Frye, A. Byron Leonard, and H. B. Willman, 1959.

Pleistocene Series

Wisconsinan Stage

Woodfordian Substage

Shelbyville Moraine

12. Till, locally slumped, calcareous where not in surface soil, gray; thickens laterally and has sand and gravel zone at base; maximum thickness 15 feet

Morton Loess

 Loess, tan to gray-tan, calcareous, meagerly fossiliferous; some color banding and rusty crenulate zones in upper part (P-573, 1 foot above base)

Farmdalian Substage

Farmdale Silt

10. Loess, pale purplish-pink, calcareous, friable; laterally variable in thickness

Altonian Substage

Roxana Silt

Zone IV

 Loess, noncalcareous, rusty brown with 1-inch dark brown streak at base, friable, massive (P-572)

Zone III

 Loess, gray, pink and rusty brown interzoned, locally organic-rich, noncalcareous, massive (P-571)

 Loess, discontinuous color band, light gray, noncalcareous, massive, (P-570)

(P-570) 0.3
6. Loess, banded pink and tan, noncal-careous, massive but with some color banding (P-569) 1.0

Zone II

 Loess, pink, noncalcareous, massive (P-568)
 1.2

Zone I

 Silt, with some sand and very few pebbles; colluvium, gray to streaked gray and rust, noncalcareous, massive (P-567) Illinoian Stage

3. Sangamon Soil; silt, sand, and a few pebbles, medium gray to dark gray, noncalcareous, friable, massive; locally faintly pelletized structure; A-zone of Sangamon Soil, but gradational with colluvium above 0.5

 Sangamon Soil; B-zone developed in till, red-brown in upper part grading downward to tans, leached; upper part has crumb to micro-blocky structure, some clay skins and a few Mn-Fe pellets
 2.5

 Glacial till, gray to blue-gray, calcareous, compact, pebbly; in upper part some zones of tan till; a few small lenses of sand and gravel

Total

1 37.3

20.0

5.0

2.5

1.5

9.0

WOLF CREEK SECTION

Measured adjacent to Interstate Highway 57, 1.9 miles south of Williamson County line, sec. 8, T. 11 S., R. 2 E., Johnson County, Illinois, 1961.

Pleistocene Series

Wisconsinan Stage

Woodfordian Substage

Peoria Loess

4. Loess, massive, gray-tan, leached, friable, locally mottled with pinktan; medium to fine silt; gray zone occurs below the B-zone of the surface soil (P-1060)

Altonian Substage

Roxana Loess

 Loess, pink-tan, gradational top and bottom, darker than above unit, massive, leached (P-1059)

Illinoian Stage

Loveland Loess

2. Loess, dark pink-brown, mealy but compact, leached, massive (P-1058)

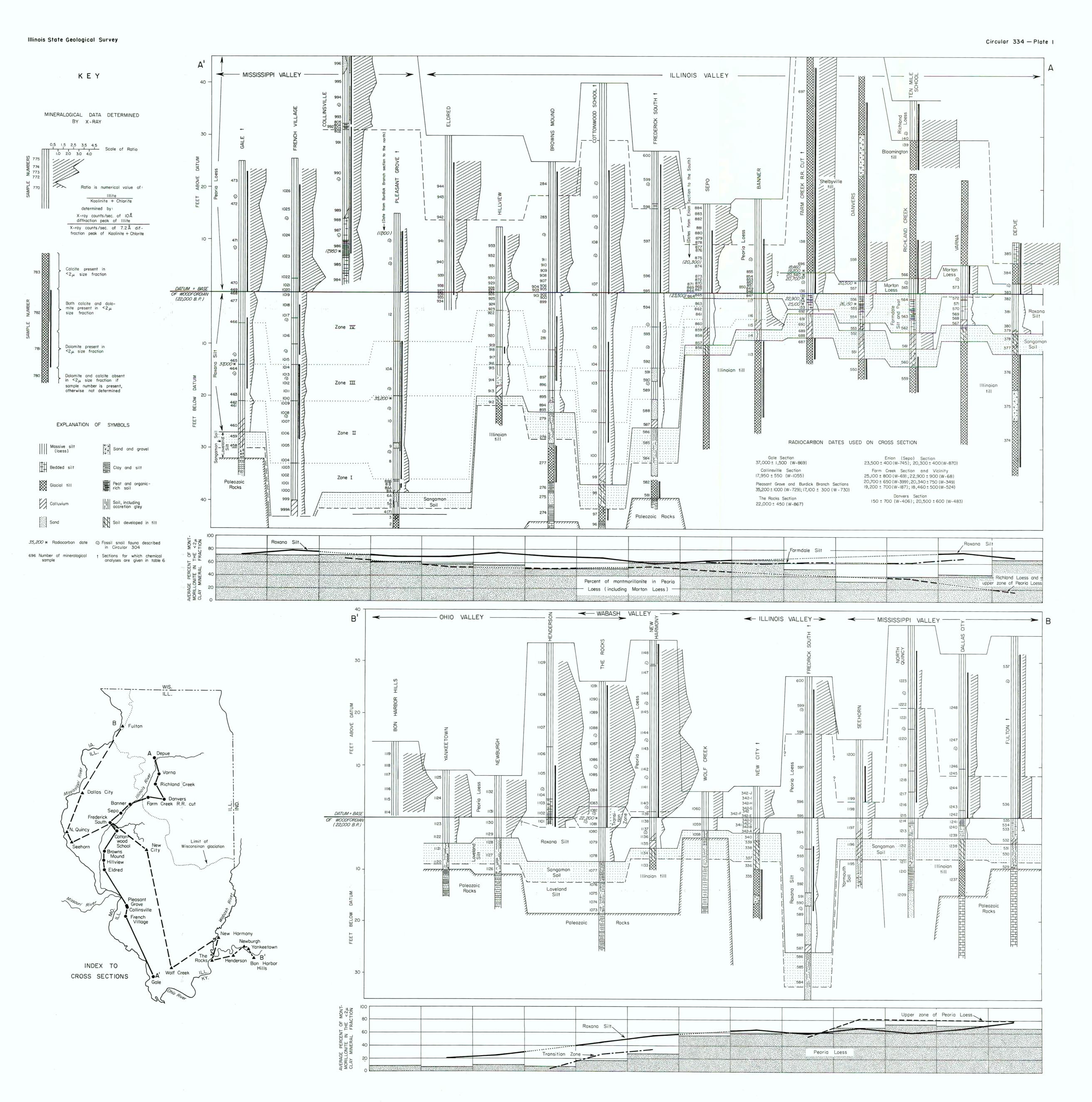
Pennsylvanian System

 Strongly developed red B-zone in Pennsylvanian sandy shale

Total Pleistocene

ickness (feet) Illinois State Geological Survey Circular 334 55 p., 1 pl., 3 figs., 6 tables, 15 geol. sections, 1962

Printed by Authority of State of Illinois, Ch. 127, IRS, Par. 58.25.



LOESS STRATIGRAPHY AND MINERALOGY OF MEASURED SECTIONS

CIRCULAR 334

ILLINOIS STATE GEOLOGICAL SURVEY URBANA